The Effect of Crater Obliteration on Inferred Surface Ages on Mars

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Overview

- Background of crater counting and obliteration
- Overview of model
- Model validation
- Results



Background

Lunar crater counts, correlated with radiometrically dated samples (and extrapolated to other bodies) remain the only quantitative tool to determine planetary surface ages





Photo credit: NASA

Unlike on the Moon, the martian atmosphere has the power to erode and deposit material, obliterating smaller craters



Previous Work

- Loss of small craters relative to lunar-derived isochons has been modeled by *Hartmann* (1971), *Chapman et al.* (1969), and observed by many others
- However, the effect of obliteration remains:
 - not commonly acknowledged in recent crater counting studies
 - poorly quantified (not tied to rates of erosion and deposition)





Simple Quantitative Model





Simple Quantitative Model





Simple Quantitative Model

Apply these inputs to a formula to describe constant production and loss:

$$N(D,t) = \frac{P(D)}{\lambda(D)} (1 - e^{-\lambda(D)t})$$

Model Inputs

P(D) – Production rate derived from Hartmann Production Function (HPF) (2005)

 $\lambda(D)$ = Loss time for a crater of diameter *D* as a function of β (= combined rates of erosion and deposition)

Assumptions

Crater loss occurs as shown (through ground lowering and infilling)
Once crater depth is reduced to zero, crater is invisible to detection
Constant production rate for last 3.5 Ga
Proportion of secondaries assumed from modeling work of *McEwen et al.* (2005)



Testing the Model

- Test model against calculated measurements of erosion rates at MER landing sites (Golombek et al., 2006)
 - Erosion rates were converted to β by assuming that the volume of eroded material is locally deposited over crater area (an assumption made by *Golombek et al.*)
- Model is fit using nonlinear regression to independent crater counts to derive β



Testing the Model



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Martian Erosion Rates





Model Results

Modeled effect of obliteration (β) on crater curves

Roll-off of crater-abundance curves at small diameters for surfaces of different ages





Model Results





Model Results



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Conclusions

- For low obliteration rates (<10 nm a⁻¹), depositional surfaces of all ages may be reliably dated, even if only small craters are counted
- For moderate to high oblit. rates (>50 nm a⁻¹), ages determined from counts of only small craters can lead to falsely low ages
- The reevaluation of a previous study demonstrates that even moderate erosion rates can significantly affect crater-count ages



Effect of secondary craters

- Secondary craters may constitute a large fraction of small craters on Mars (*McEwen et al.*, 2005)
- HPF, which our model is based upon, includes both secondaries and primaries (*Hartmann*, 2005 & 2007), though:
 - the distribution of secondaries may be spatially heterogeneous, with large populations near large primary impacts and associated rays
 - a martian impact may produce more secondaries than a lunar one
- Both effects would locally <u>increase</u> populations of small craters and our model-derived values for obliteration (β) may be taken as minima



Model Inputs

 $P(D) = 0.29 \begin{cases} \frac{0.0035(0.13\ln(D) + 0.83)}{d^{3.3}} & D > 0.\\ 10^{-1.8\log(D) - 2.59} & D \ge 1.\\ 10^{-2.2\log(D) - 1.89} & D \ge 1. \end{cases}$

D > 0.001 and D < 1.4 $D \ge 1.4 and D \le 48.1$ D > 48.1

For known input crater population (primary or secondary)

$$\lambda(D) = \frac{\beta}{1000\xi}; \quad \xi = \begin{cases} 0.2 \cdot D_p \text{ or } 0.1 \cdot D_s & D < 5.8\\ 0.42 \cdot \ln(D) - 0.01 & D \ge 5.8 \end{cases}$$

$$N(D,t) = \frac{1000\xi P(D)}{\beta} \left(1 - e^{-\beta t/(1000\xi)}\right);$$

$$\xi = \begin{cases} 0.2 \cdot D_p & D < 5.8\\ 0.42 \cdot \ln(D) - 0.01 & D \ge 5.8 \end{cases}$$

For unknown crater population (mixed primaries/secondaries)

$$\lambda(D) = \frac{\psi\beta}{1000\xi}$$

$$\boldsymbol{\chi} = \begin{cases} 1 + \prod_{s} & D < 1.2 \\ 0 & D \ge 1.2 \end{cases}; \quad \boldsymbol{\xi} = \begin{cases} 0.2D & D < 5.8 \\ 0.42 \ln(D) - 0.01 & D \ge 5.8 \end{cases}$$

$$N = \frac{1000}{\psi\beta} P(D) \xi \left(1 - e^{-\frac{\psi\beta t}{1000\xi}} \right)$$



Derivation of model formula

$$\begin{aligned} \frac{dN}{dt} &= P - \lambda \cdot N \\ \frac{dN}{dt} &= P \left(1 - \frac{\lambda}{P} \cdot N \right) \\ \frac{dN}{\left(1 - \frac{\lambda}{P} \cdot N \right)} &= P \cdot dt \\ \int \frac{\left(-\frac{P}{\lambda} \cdot d\Gamma \right)}{\Gamma} &= fP \cdot dt ; \\ \Gamma &= 1 - \frac{\lambda}{P} \cdot N , \ d\Gamma &= -\frac{\lambda}{P} \cdot dN \\ - \frac{P}{\lambda} \cdot \ln(\Gamma) &= P \cdot t \\ \ln \left(1 - \frac{\lambda}{P} \cdot N \right) &= -\lambda \cdot t \\ e^{-\lambda \cdot t} &= 1 - \frac{\lambda}{P} \cdot N \\ N &= \frac{P}{\lambda} \cdot (1 - e^{-\lambda \cdot t}) \end{aligned}$$

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Mid-Latitude Crater in Noachis Terra (14.5E, 49.2S)

- Crater floor fills with wind-blown fines and boulders falling from rim
- Rim and ejecta blanket are largely lost to erosion
- Crater rim becomes dissected and rounded as material sheds onto floor





Syria Planum (269W, 32.47S)

Crater near detection limit has no visible contrast with surrounding material



MOC-NA S10-01620



500 m

MOC-NA S06-00539

Protei Regio (south of Capri Chasma) (53.85W, 19.34S)

Craters in varying degrees of degradation; no circular forms without topographic expression

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Pedestal Craters

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Northern Plains (246.96W, 61.5N)





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100 m



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MOC R13-02372

1 km

HIRISE PSP_001358_2485



S of Noctis Labyrinthus (95.57W, 15.44S)

• Crater rim composition is visibly different and emphasizes craters for identification

• Degradation removes lightcolored rim and lessens contrast

• Light-colored rim rock combines to make a lighter crater fill

• Some vaguely circular light – toned patches in plains, but mostly irregular and difficult to count

500 m

MOC S08-01945



Production rate assumption

 Constant production rate

 Positive factor of error 1.6
 Falls within error of HPF (Factor of 2-3 introduced by R_{bolide})