

Lecture 17 – Mars spectroscopy

Reading

Ch 4.7 Forest applications

What was covered in the previous lecture

LECTURES

- Jan 05 1. Intro
- Jan 07 2. Images
- Jan 12 3. Photointerpretation
- Jan 14 4. Color theory
- Jan 19 5. Radiative transfer
- Jan 21 6. Atmospheric scattering
- Jan 26 7. Lambert's Law
- Jan 28 8. Volume interactions
- Feb 02 9. Spectroscopy
- Feb 04 10. Satellites & Review
- Feb 09 11. **Midterm**
- Feb 11 12. Image processing
- Feb 16 13. Spectral mixing
- Feb 18 14. Classification
- Feb 23 15. Radar & Lidar
- **Feb 25 16. Thermal infrared previous**
- **Mar 02 17. Mars spectroscopy (Matt Smith) today**
- Mar 04 18. Forest remote sensing (Van Kane)
- Mar 09 19. Thermal modeling (Iryna Danilina)
- Mar 11 20. Review
- Mar 16 21. **Final Exam**

Friday's lecture

Forest remote sensing

Today's lecture:

Mars spectroscopy

Next lecture – Forest remote sensing (Guest: Dr. Van Kane, CFR)



GEOLOGIC HISTORY OF QUARTZ-BEARING DEPOSITS IN SYRTIS MAJOR, MARS

Matt Smith
University of Washington
ESS 421
March 2, 2011

MARS SPECTROSCOPY

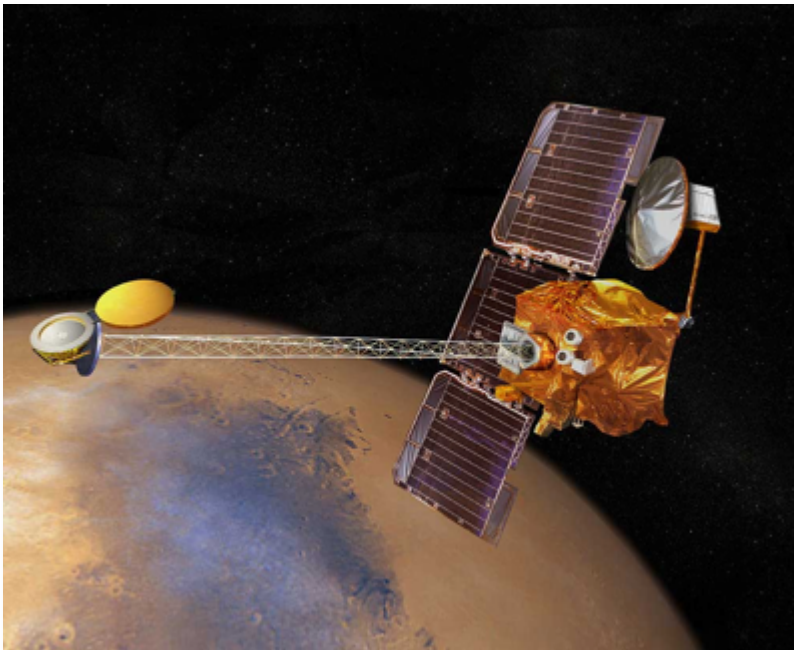
- Rovers and satellites are the most common ways that we can observe the Martian surface

Rovers offer:

Very detailed geology

Chemistry, mineralogy, and imagery of rocks and soils

Limited mobility and range



Satellites offer:

Wide spatial coverage (regional and global)

Context for rover observations

Limited resolution and detection capability

MARTIAN ORBITAL SPECTROMETERS

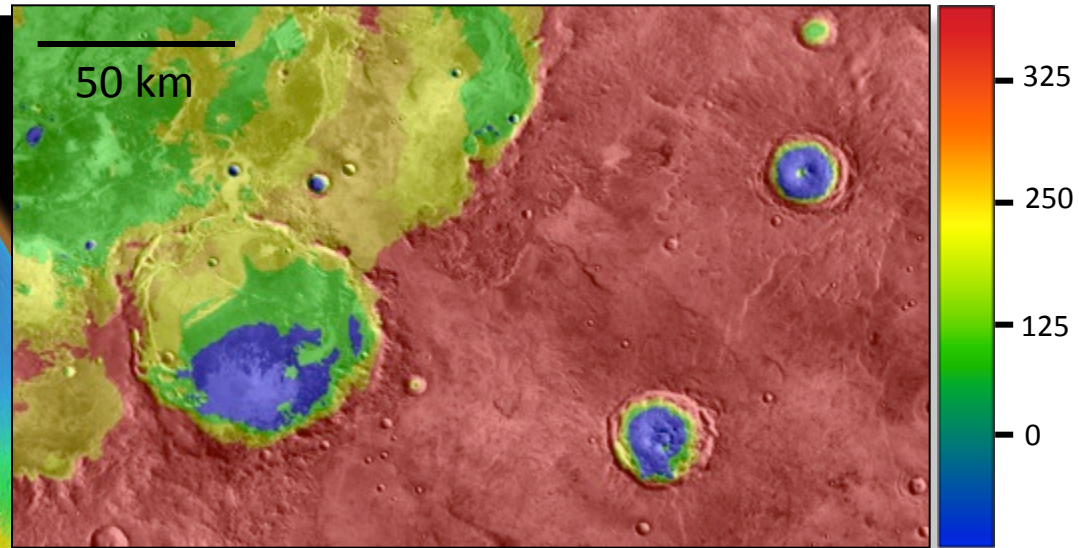
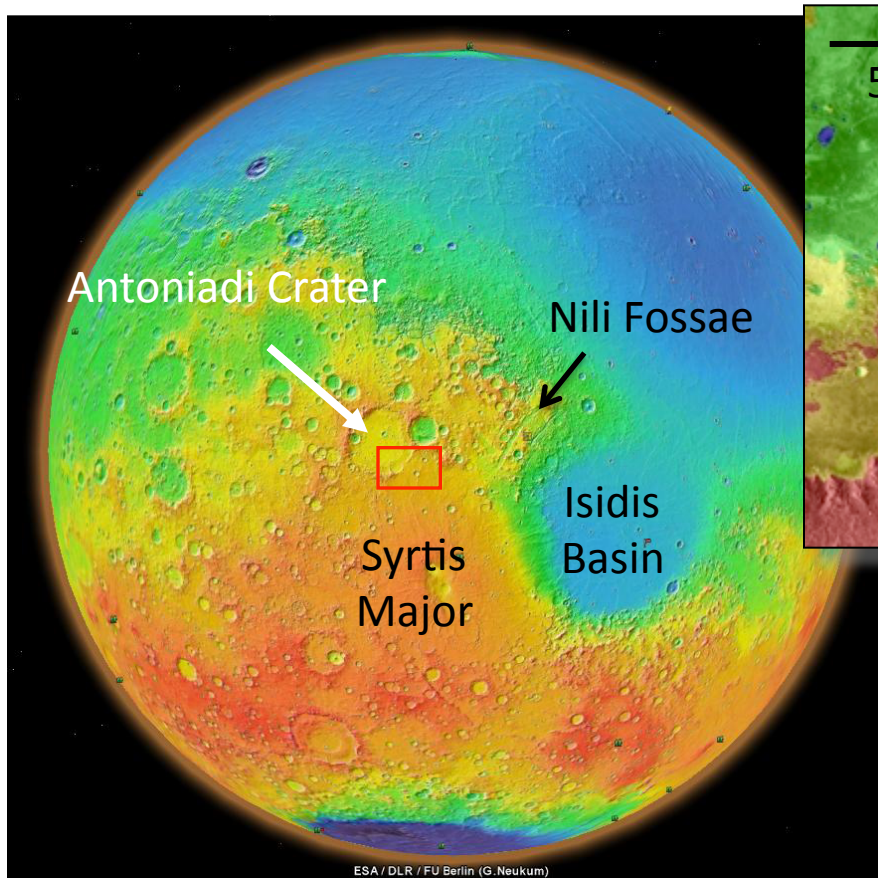
VISIBLE-NEAR INFRARED

- *CRISM* (COMPACT RECONNAISSANCE IMAGING SPECTROMETER FOR MARS)
 - 0.36 – 3.92 μm range, 16-20 m/pixel
 - Good for detection hydrated and iron-bearing minerals

THERMAL

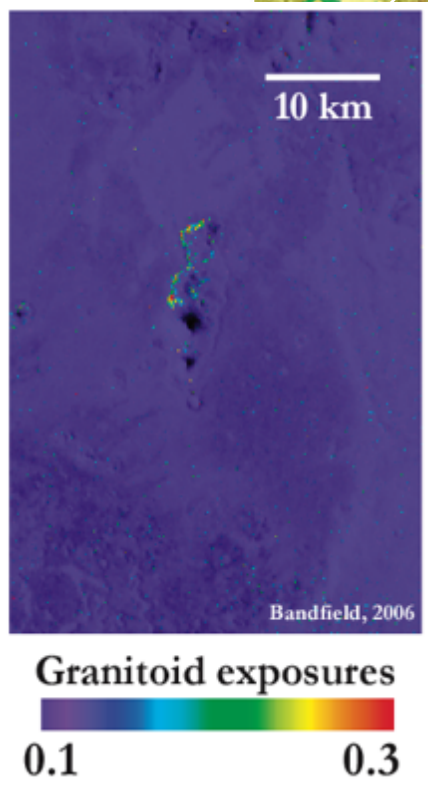
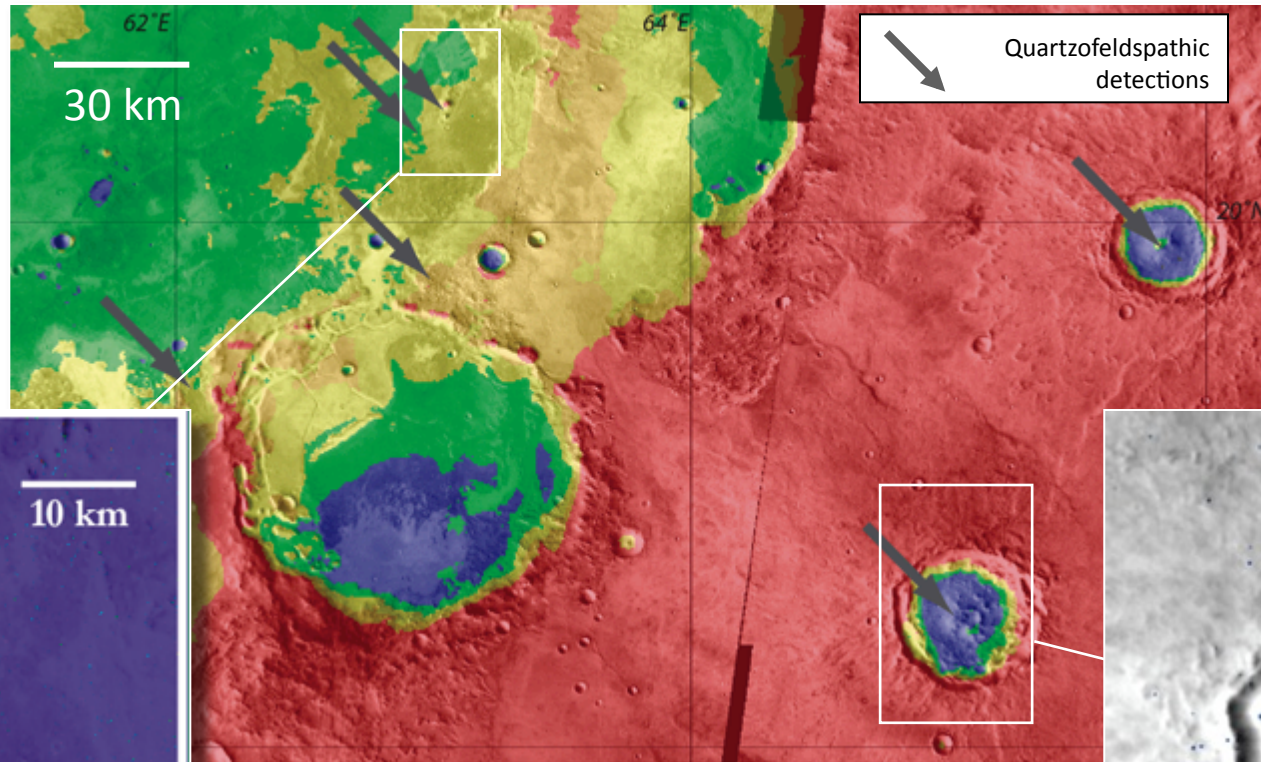
- *TES* (THERMAL EMISSION SPECTROMETER)
 - 6-50 mm range, 3 km/pixel
 - Good for detecting bulk mineralogy (i.e. basalt vs. dust)
- *THERMIS* (THERMAL EMISSION IMAGING SYSTEM)
 - 10 infrared spectral bands between 6-15 mm, 100 m/pixel
 - Used to detect large mineralogical differences between units

ANTONIADI CRATER

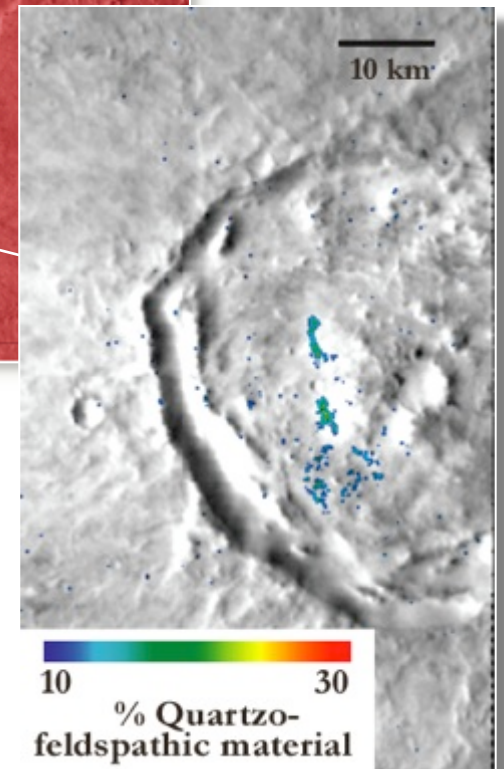


- The Antoniadi Crater region has evidence of:
- Stable long-term water on ancient Mars
 - Past volcanic activity
 - Minerals (clays, quartz/silica) that indicate a hospitable and stable environment for life

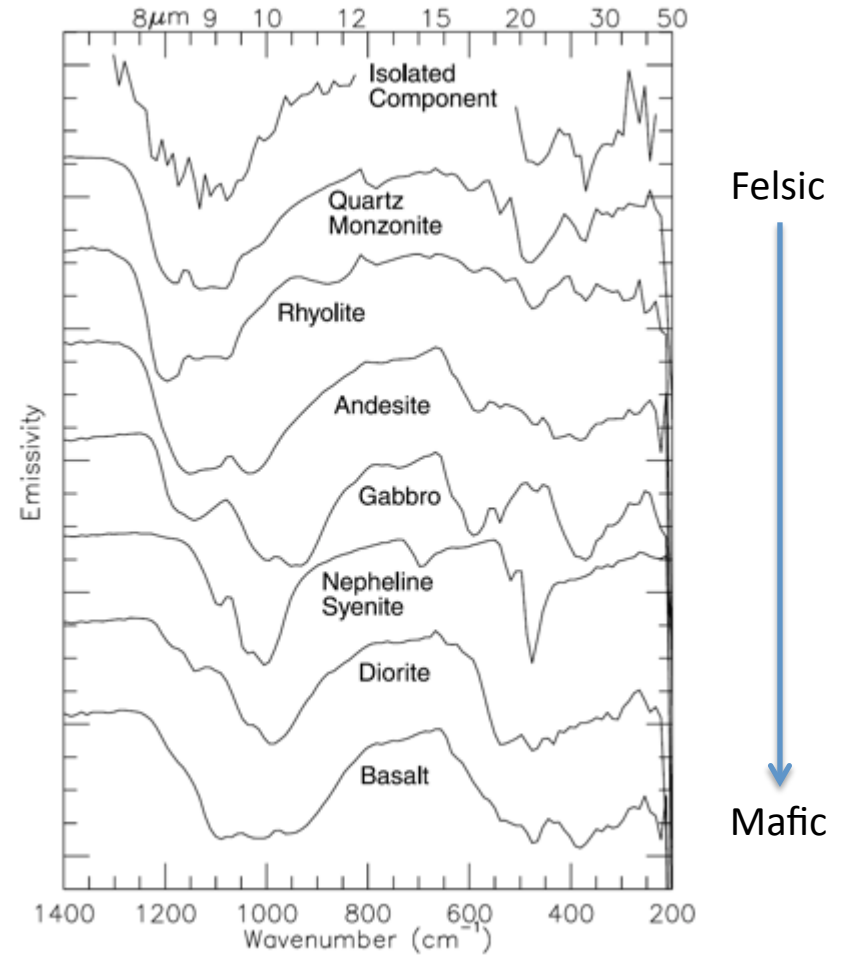
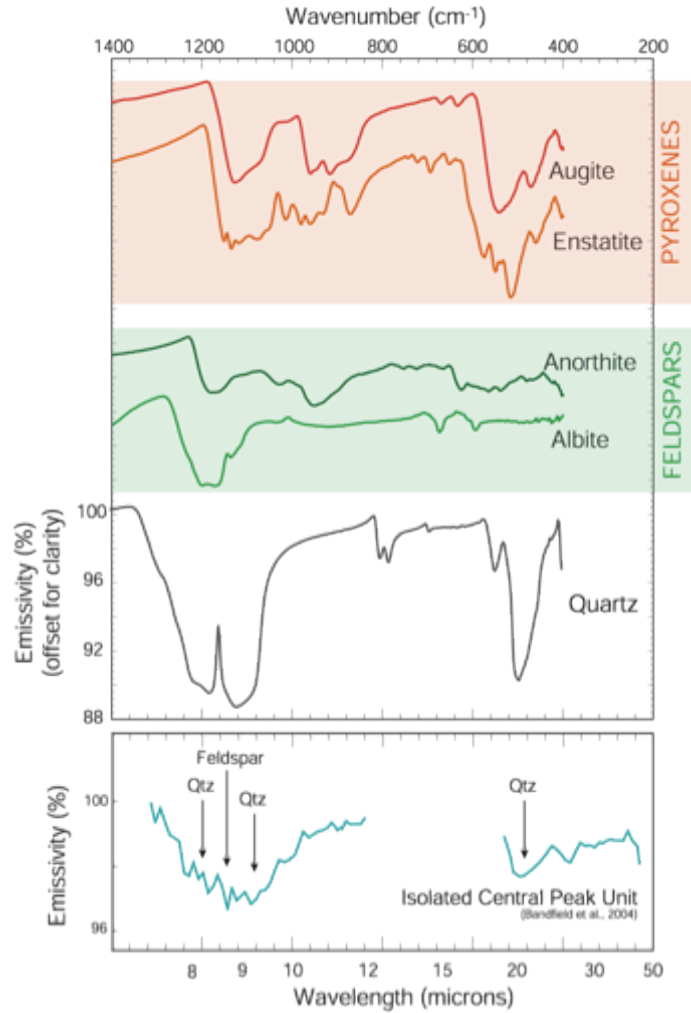
QUARTZOFELDSPATHIC UNITS



*Quartzofeldspathic units –
Rocks high in quartz and feldspar,
low in pyroxenes*

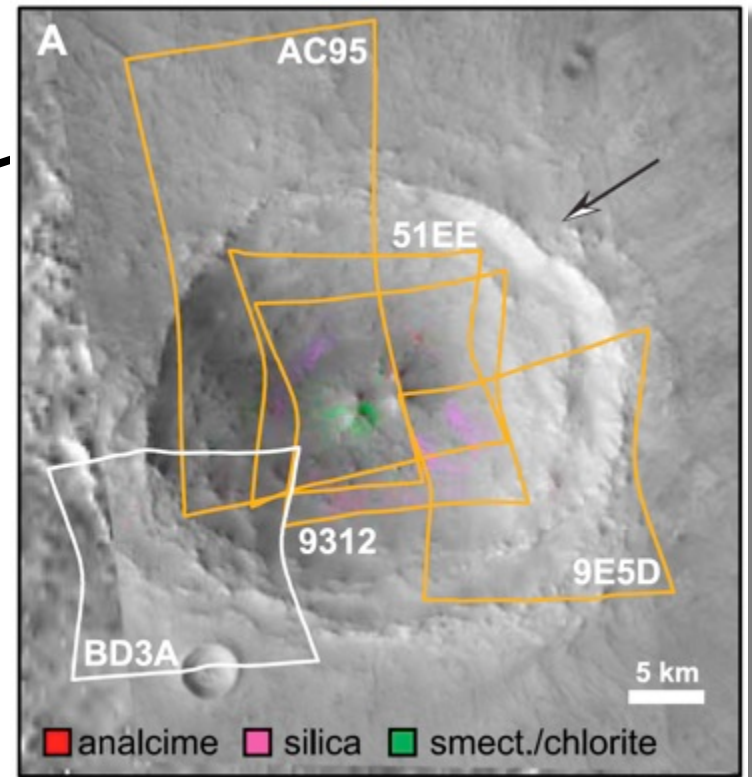
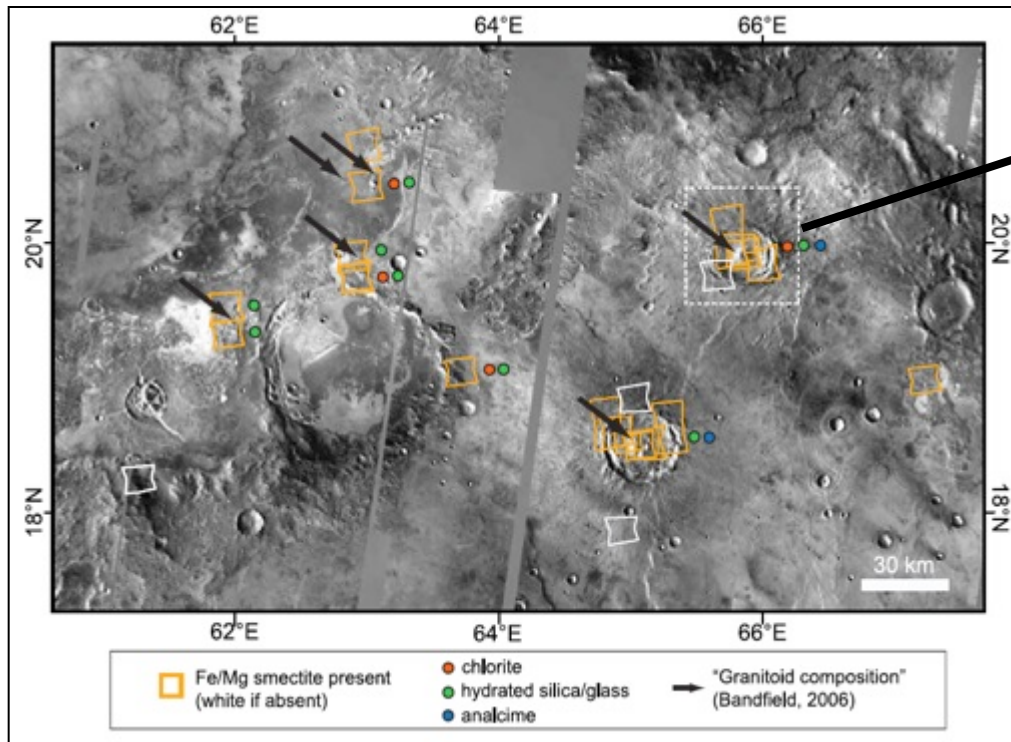


QUARTZOFELDSPATHIC DETECTION



Bandfield et al., 2004

HYDRATED MINERALS



Ehlmann et al., 2009

Identification of hydrous alteration products with CRISM:
hydrated silica, phyllosilicates (smectite, chlorite), and zeolites (analcime)

MOTIVATION FOR STUDYING ANTONIADI CRATER

Locally

- Why do we find these minerals (quartzofeldspathic unit, hydrated silica and phyllosilicates) together?
- What does their location imply for past water, temperature, or life?

Regionally

- How do these deposits correspond to nearby Nili Fossae?

Globally

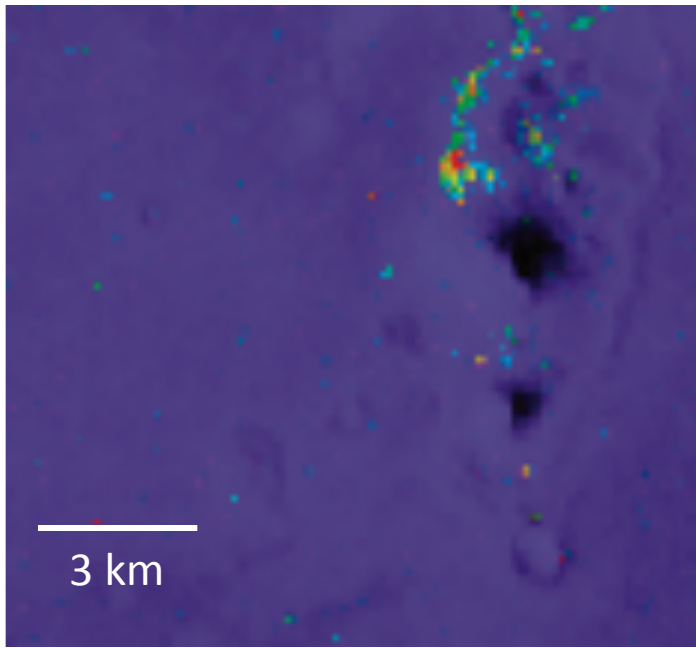
- Are these deposits similar to other mineral detections on the planet?

DESCRIPTION OF EXPOSURES:

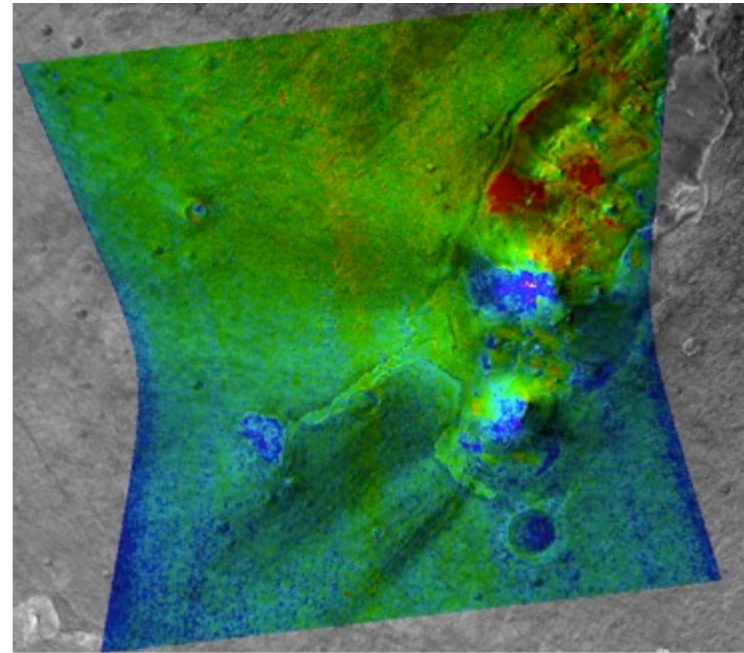
QUARTZOFELDSPATHIC/HYDRATED SILICA UNITS

- All exposures of quartzofeldspathic material are found coincident with hydrated silica
- Silica-bearing units are in thin, mobile, dune-forming deposits in topographic lows
- Intensity of detection *increases* away from source, suggesting a lag deposit
- In shallowest fractures, only hydrated silica (no phyllosilicates) is detected, indicating shallower burial

SILICA-BEARING UNITS



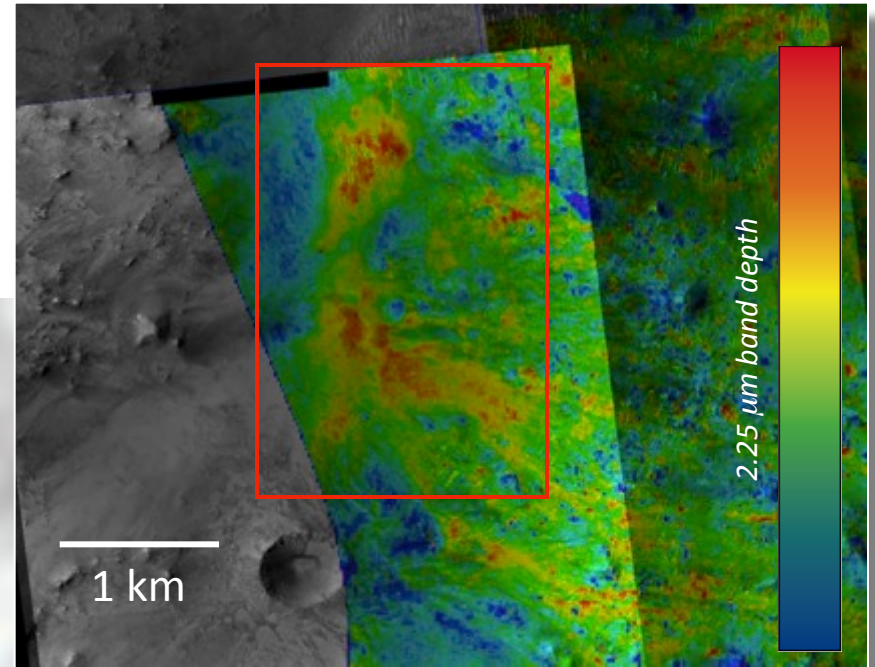
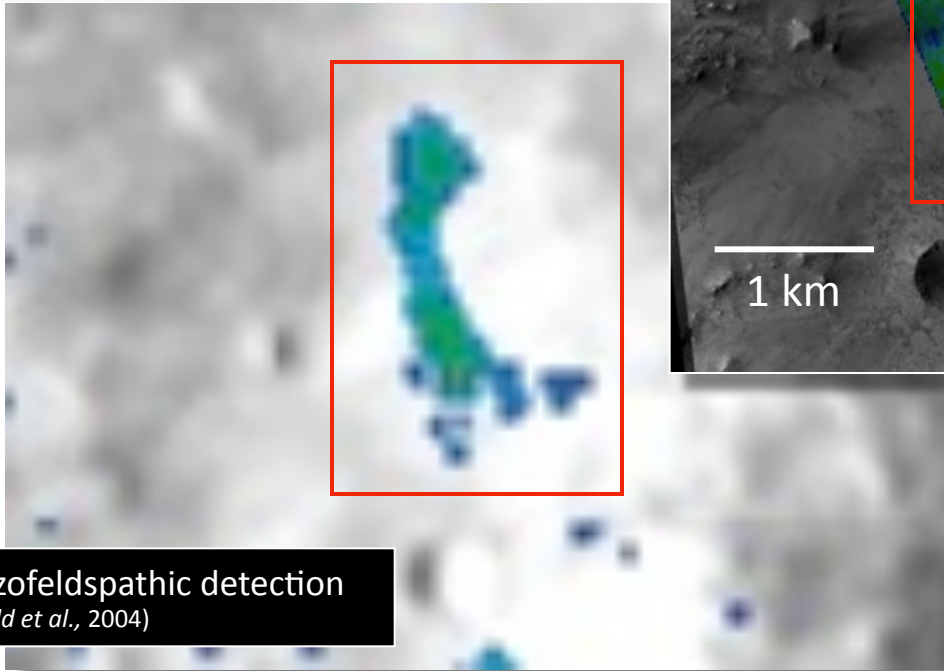
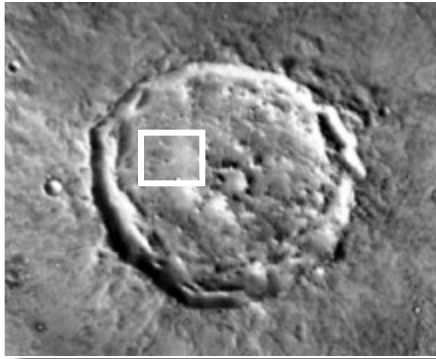
Quartzofeldspathic detection
(*Bandfield et al., 2004*)



Hydrated silica detection

Quartzofeldspathic and hydrated silica units are co-located

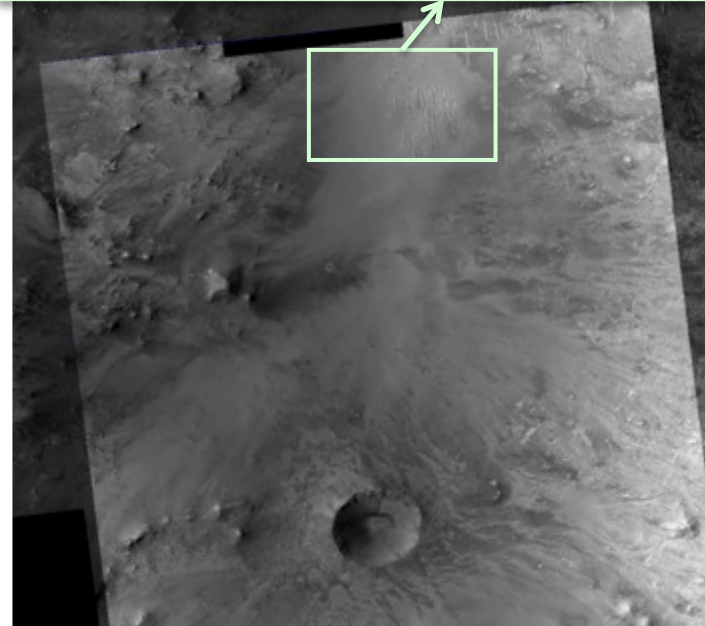
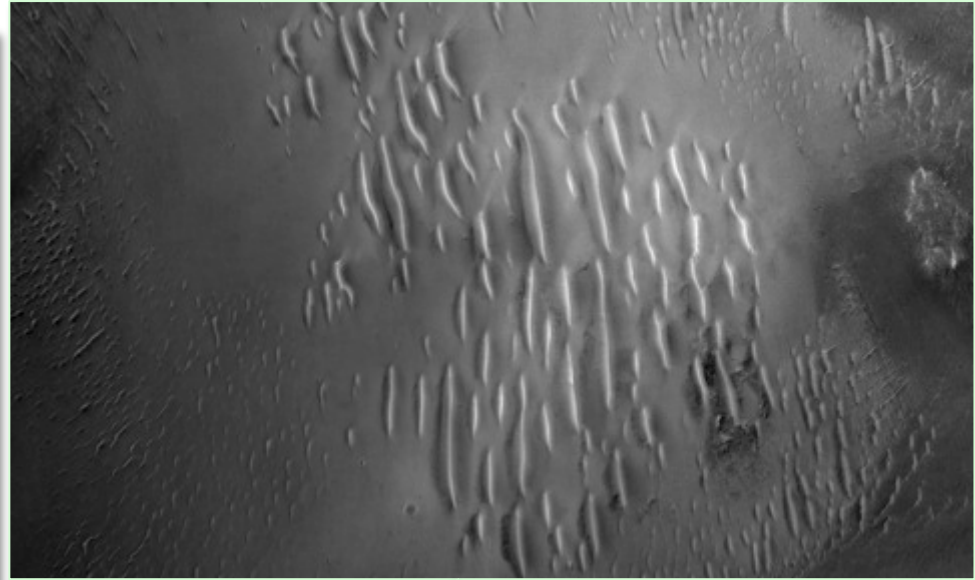
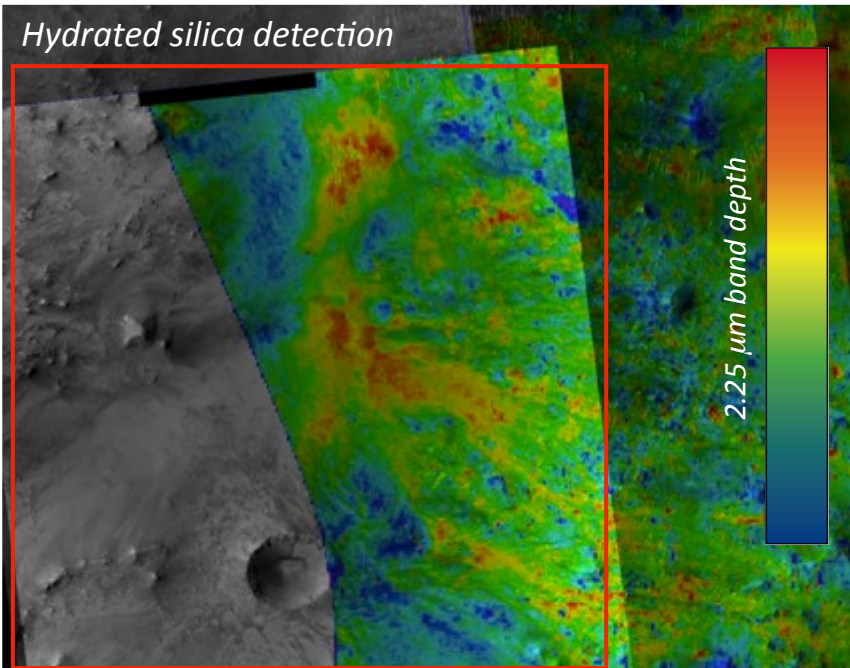
SILICA-BEARING UNITS



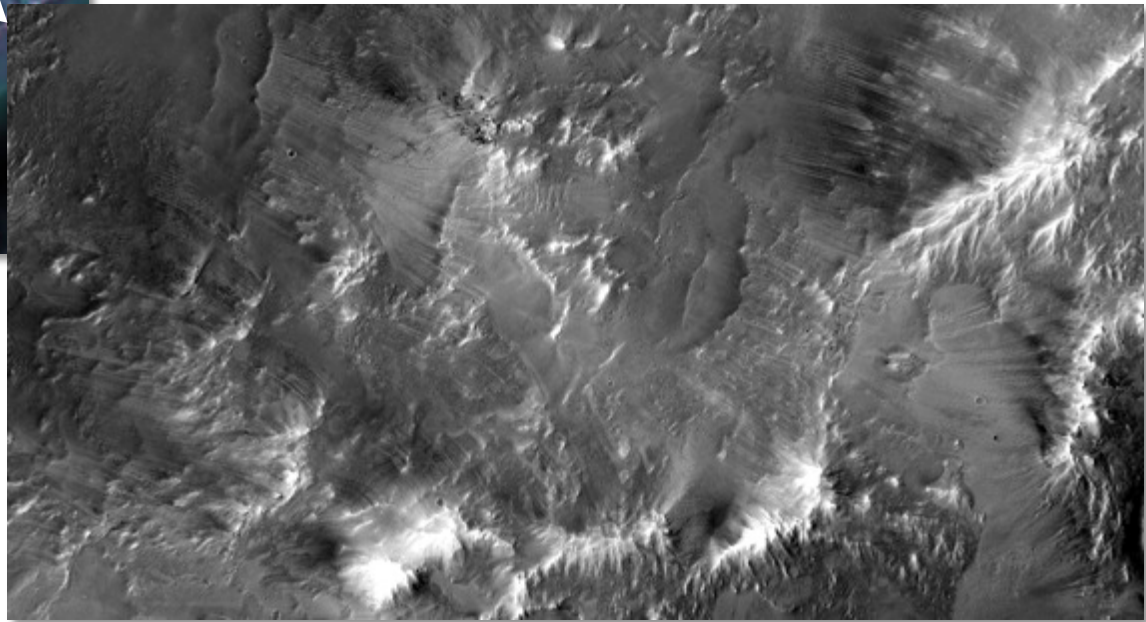
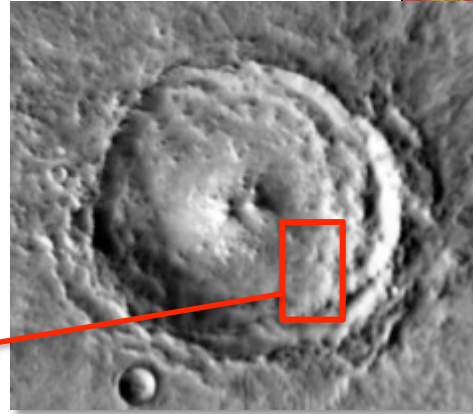
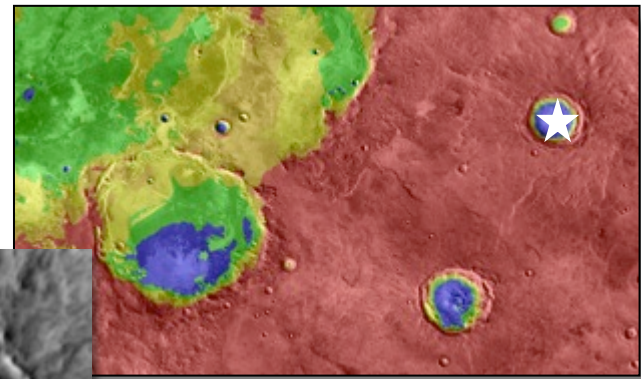
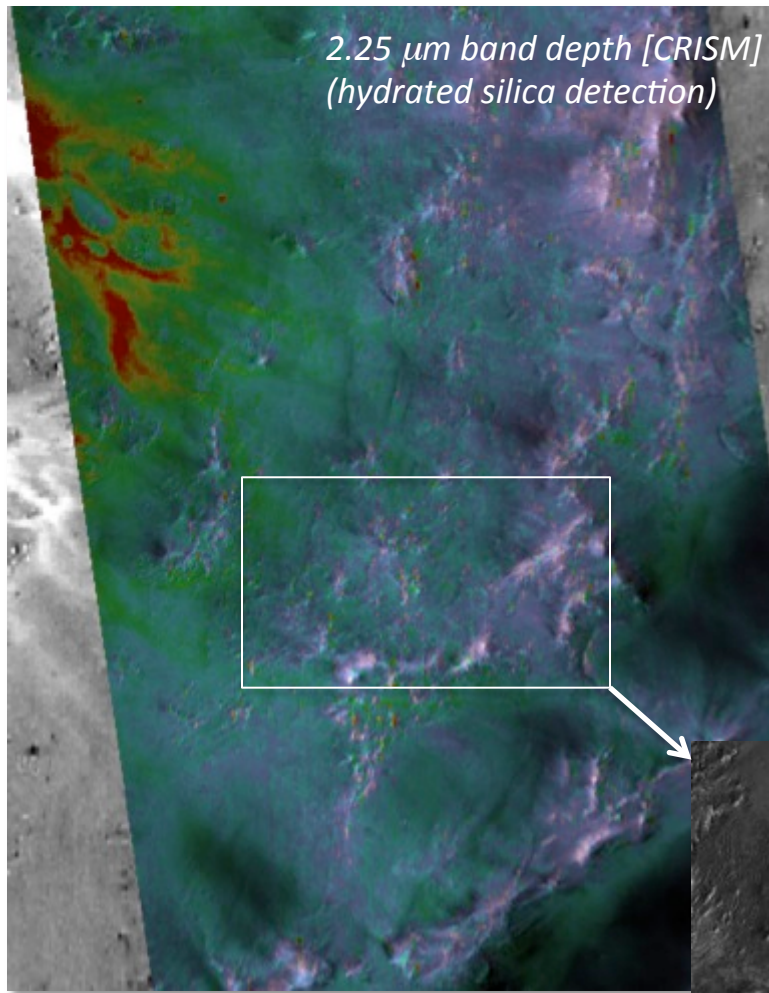
Hydrated silica detection

Quartzofeldspathic and hydrated silica units are co-located

SILICA-BEARING UNITS

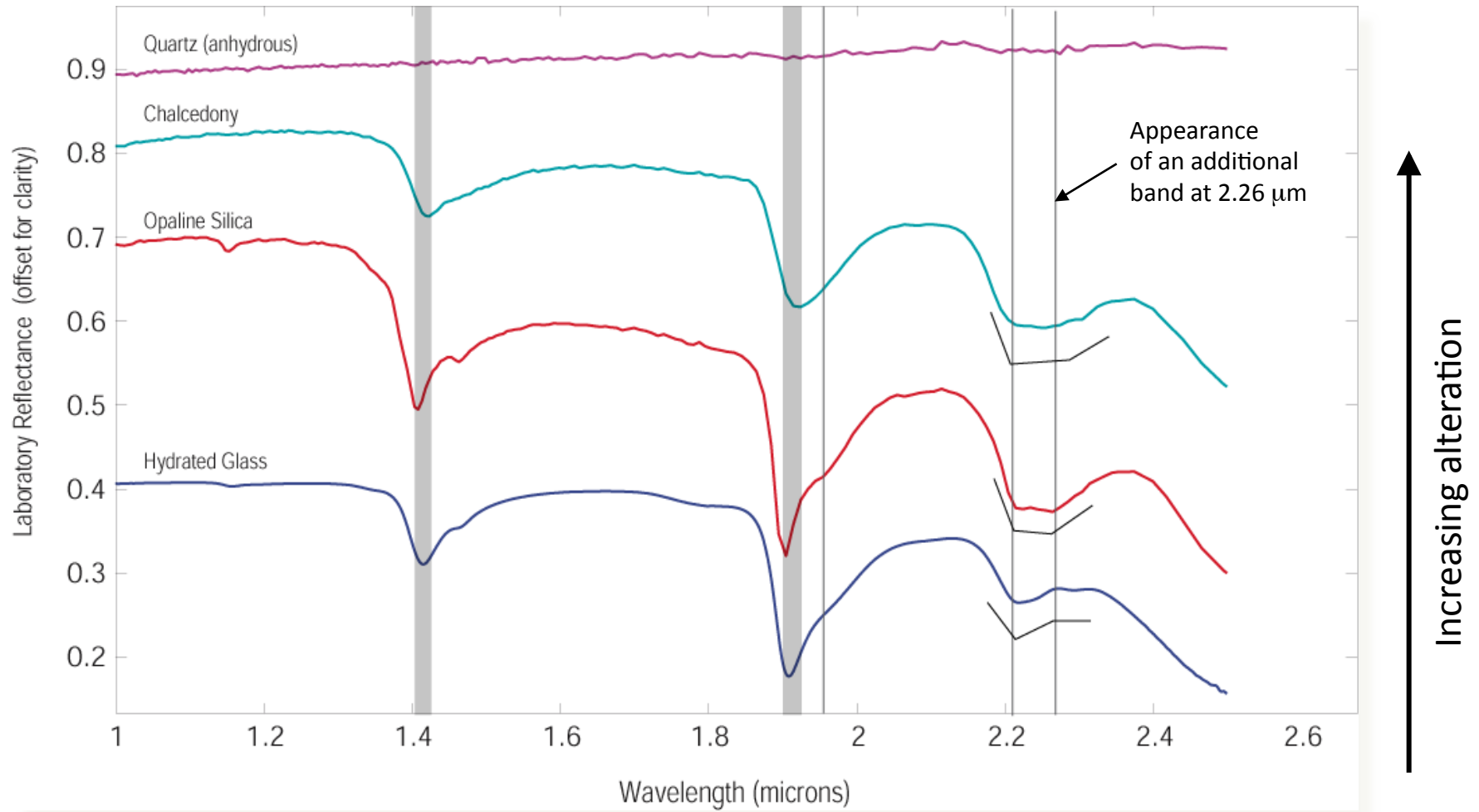


Hydrated silica/quartzofeldspathic units are associated with mobile, dune-forming units



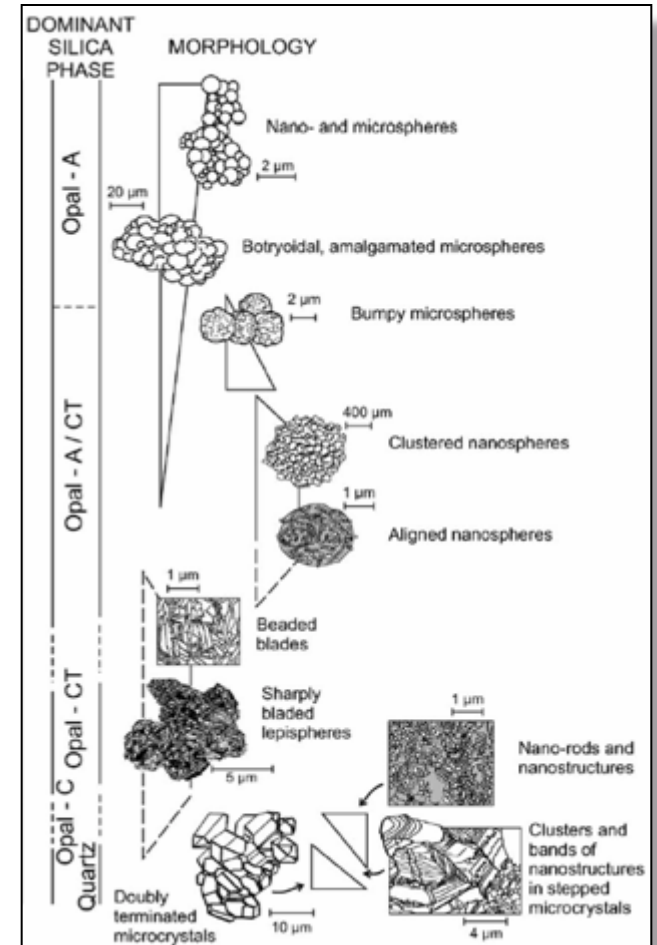
There is no obvious source stratum, and silica detection intensifies downslope, suggesting that this is a **lag deposit**

HYDRATED SILICA

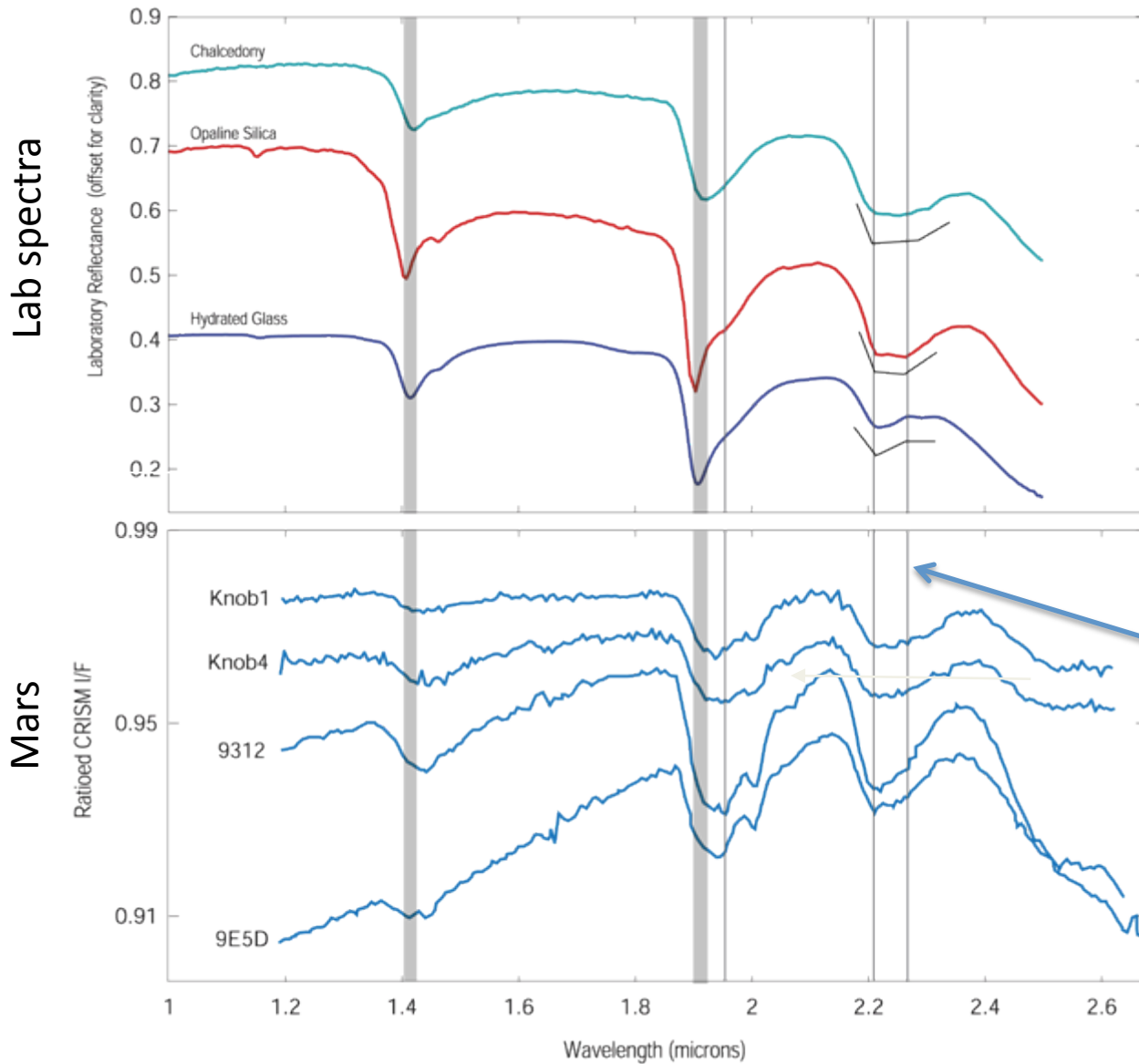


HYDRATED SILICA

- Forms under low temperature (<150°C) conditions on Earth, and typically found in near-surface environments (*Roberts et al., 1974*)
- Alters to microcrystalline quartz (chalcedony) in the presence of water (*Lynne et al., 2005*) in relatively short time scales; ~400 Ma (*Tosca and Knoll, 2009*)



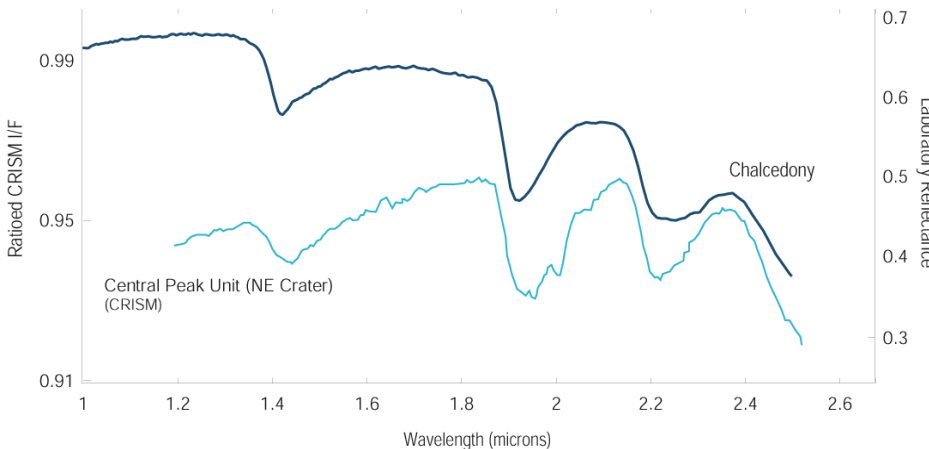
HYDRATED SILICA



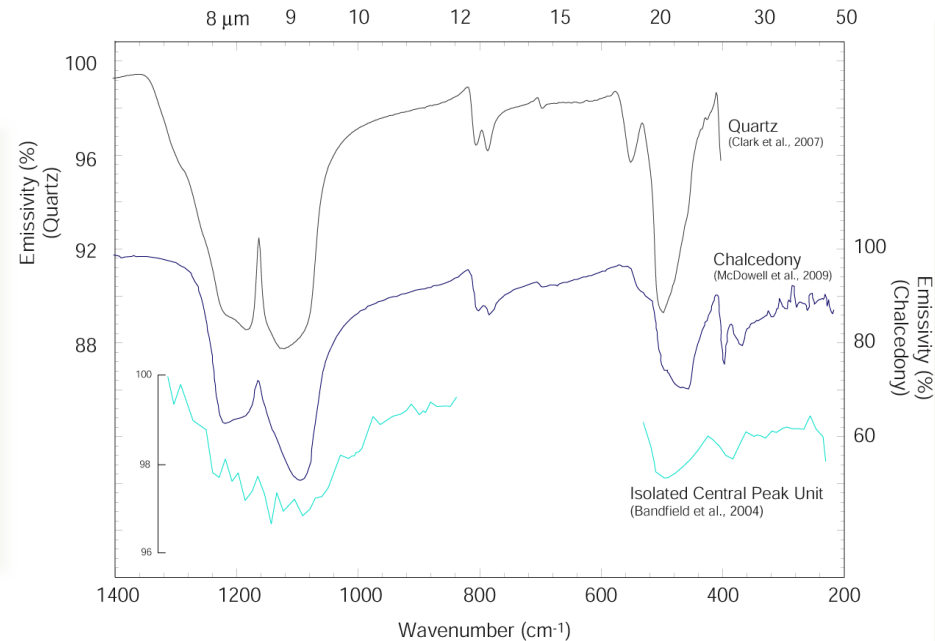
2.26 μm band indicates the presence of chalcedony (more heavily altered)

QUARTZ + HYDRATED SILICA = CHALCEDONY

Near-infrared



Thermal



Chalcedony indicates higher degree of alteration than opaline silica and locally *sustained* water (up to 400 Ma) in moderate (<150°C) temperatures

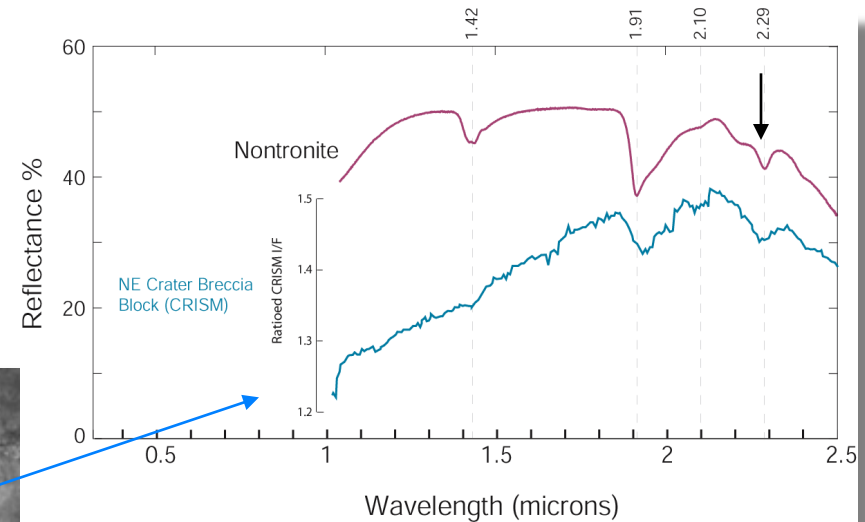
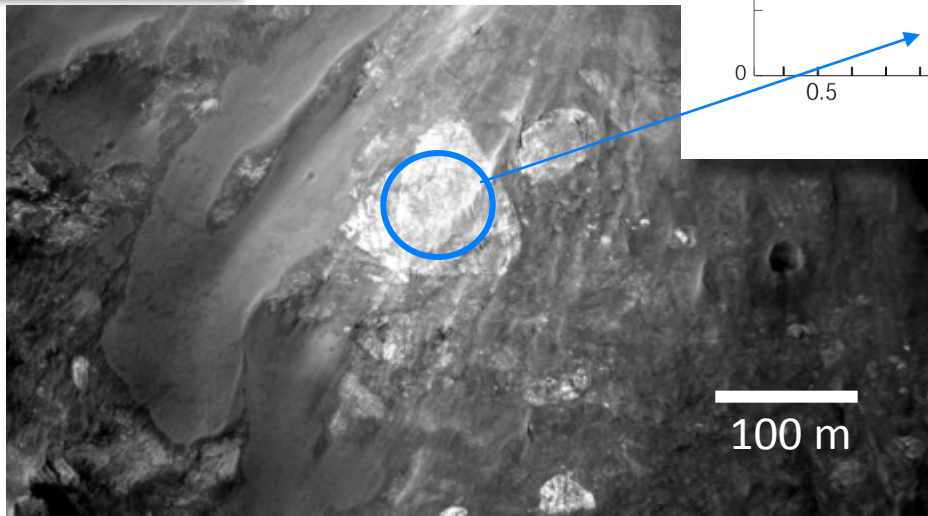
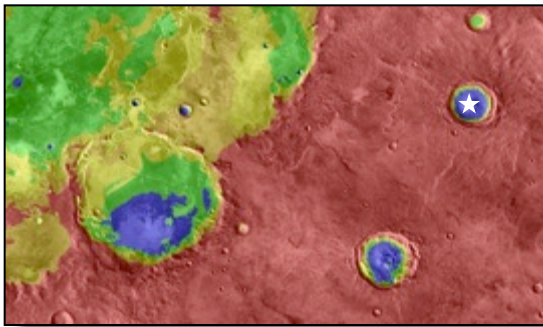
DESCRIPTION OF EXPOSURES:

PHYLLOSILICATES

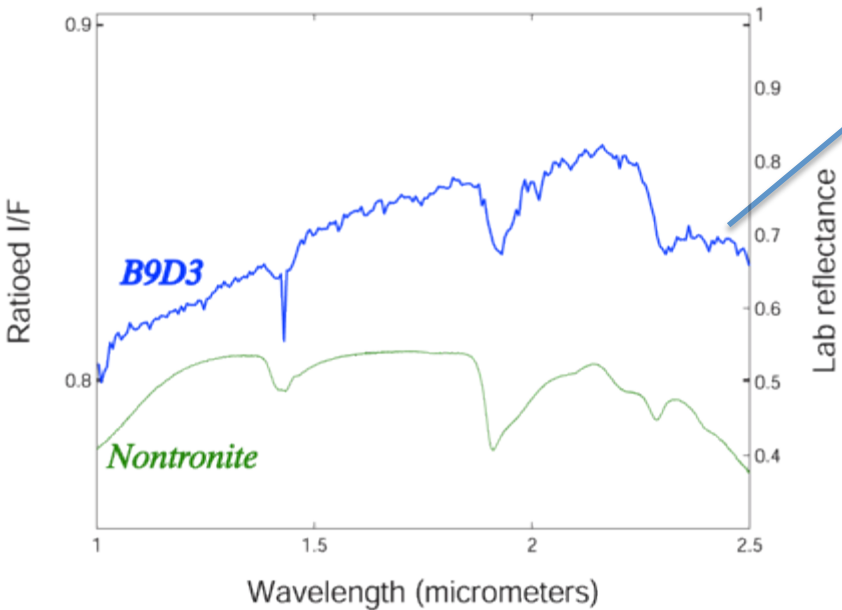
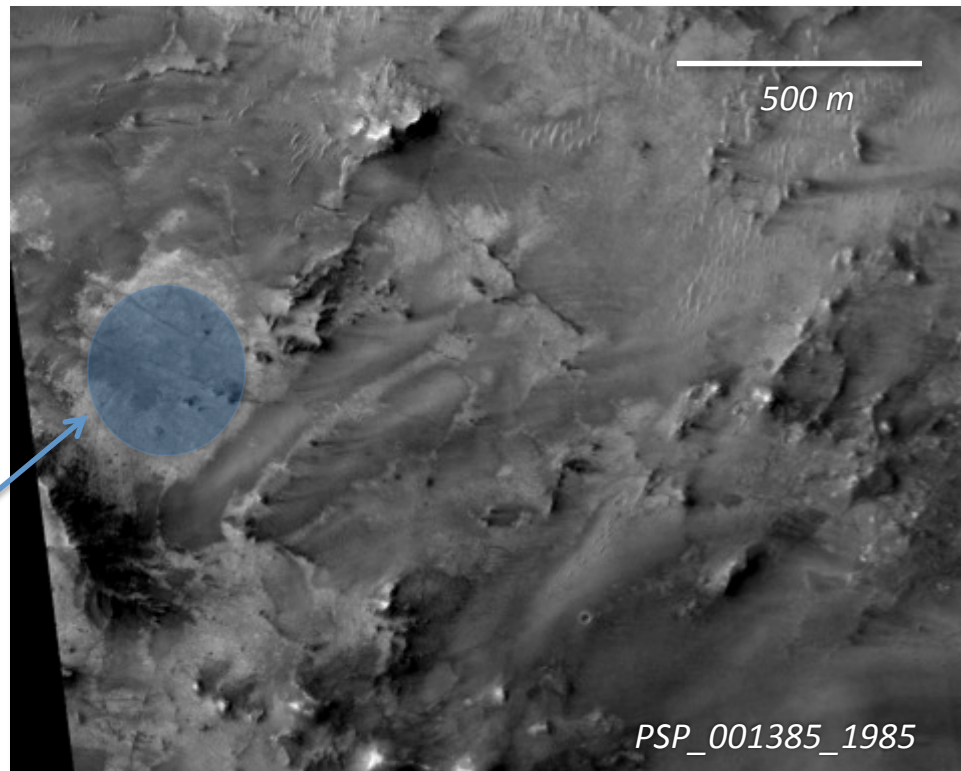
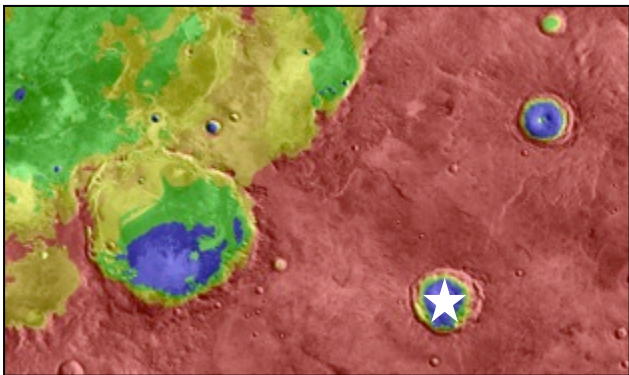
- CRISM detections of phyllosilicates are found mostly in megabreccia blocks
- Clays also associated with layered knobs (possibly ancient layered impact breccias)
- Detections of phyllosilicates are **not** mixed with hydrated silica

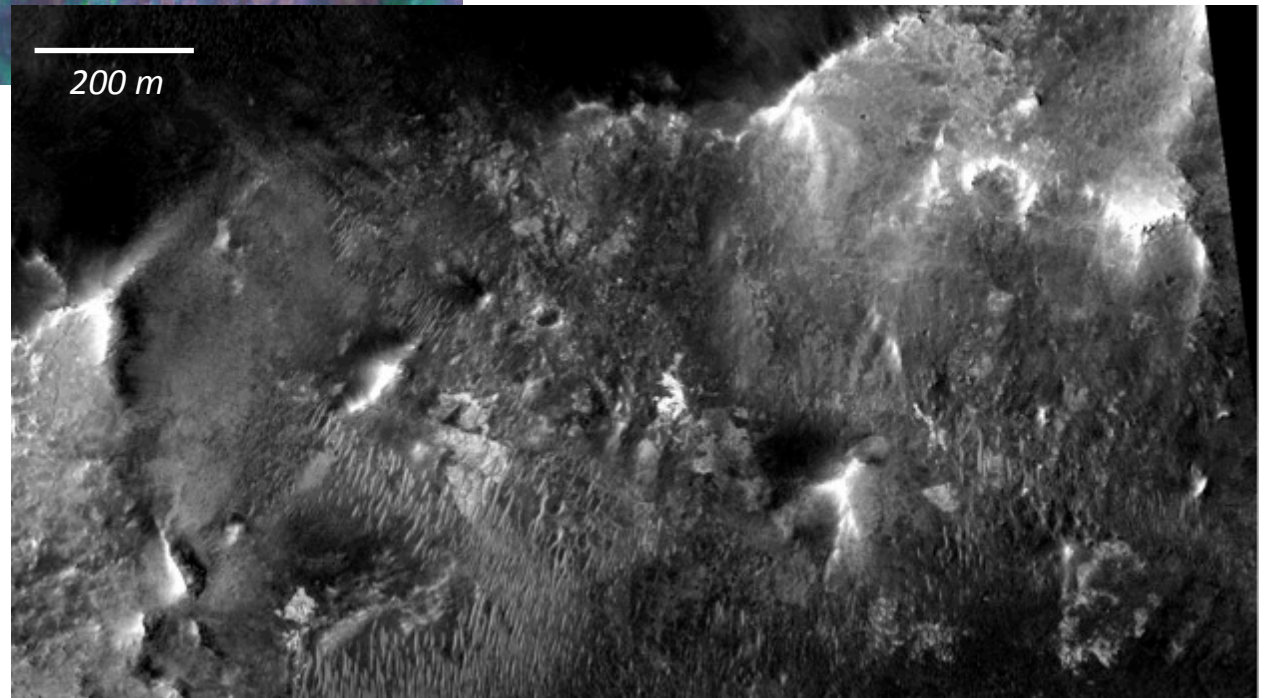
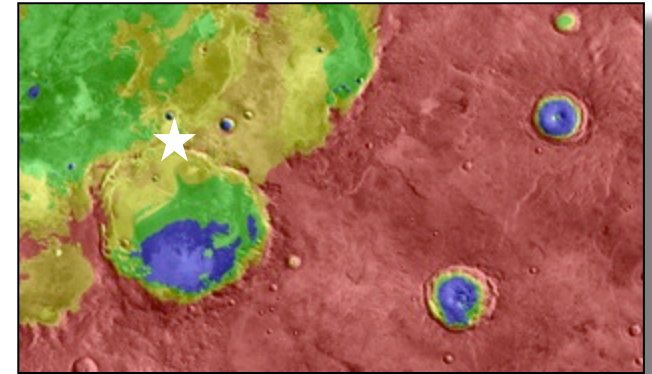
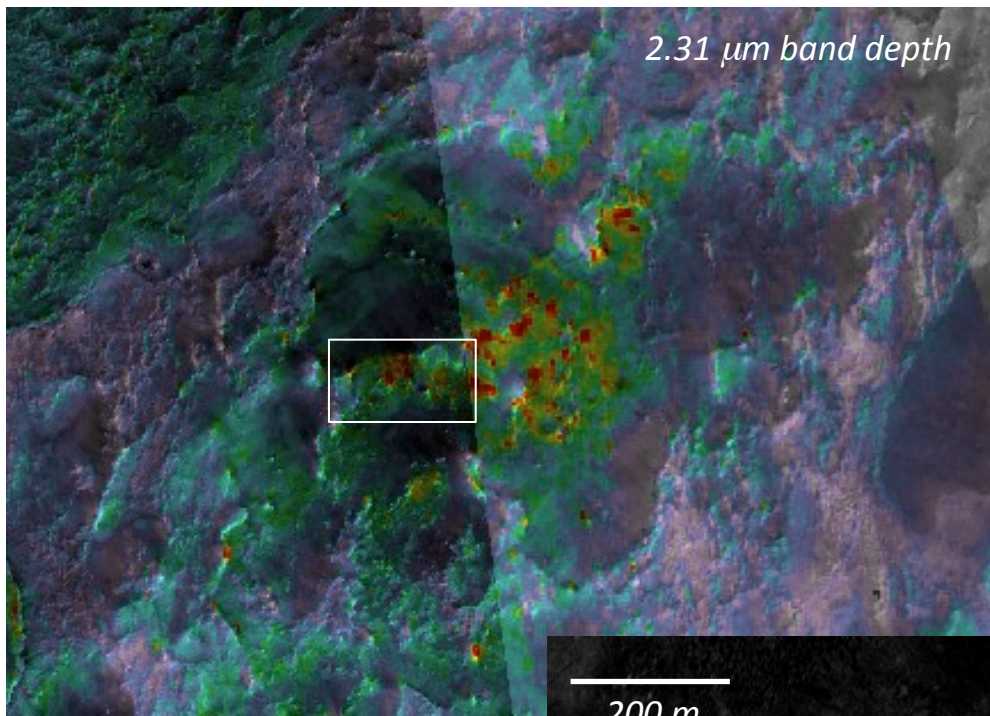
PHYLLOSILICATES

- Most exposures are in impact breccia clasts in crater central peaks and floors



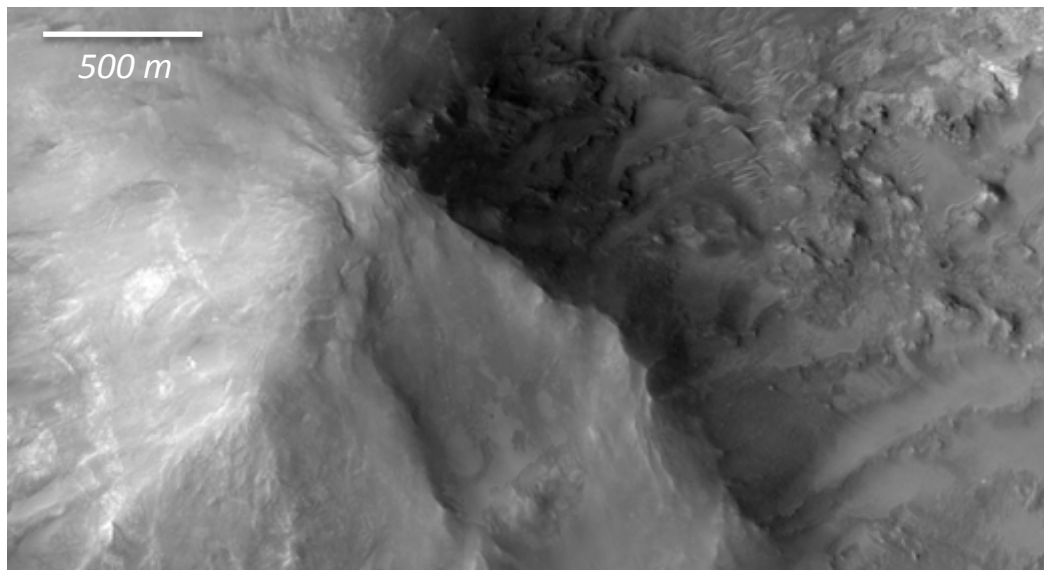
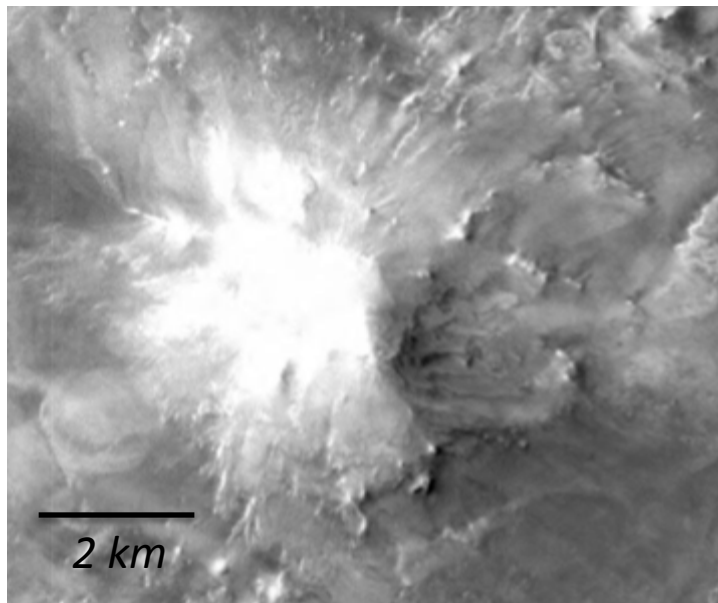
Indicates that alteration *pre-dates* and is exposed by the impact event



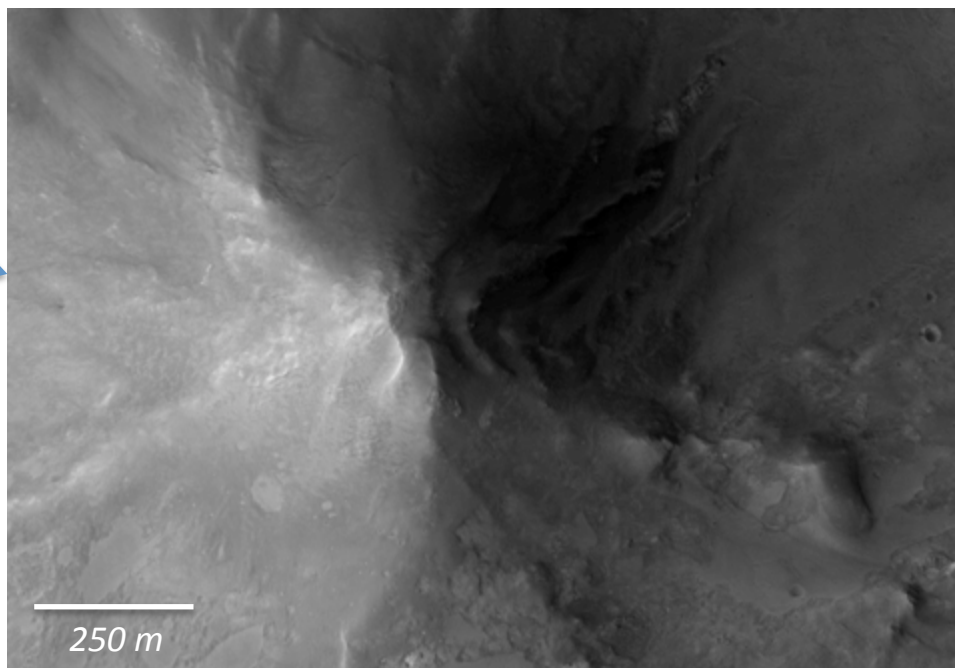
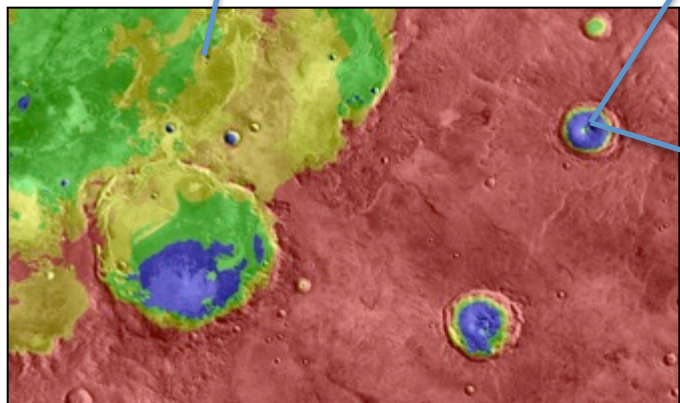


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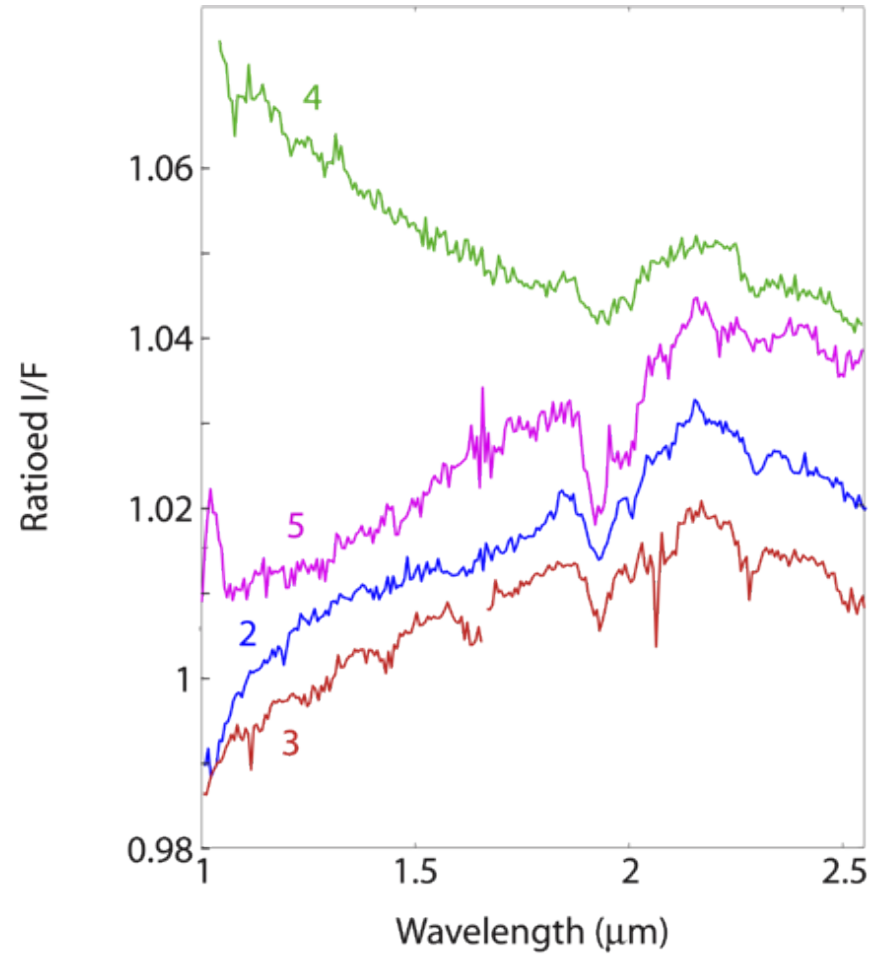
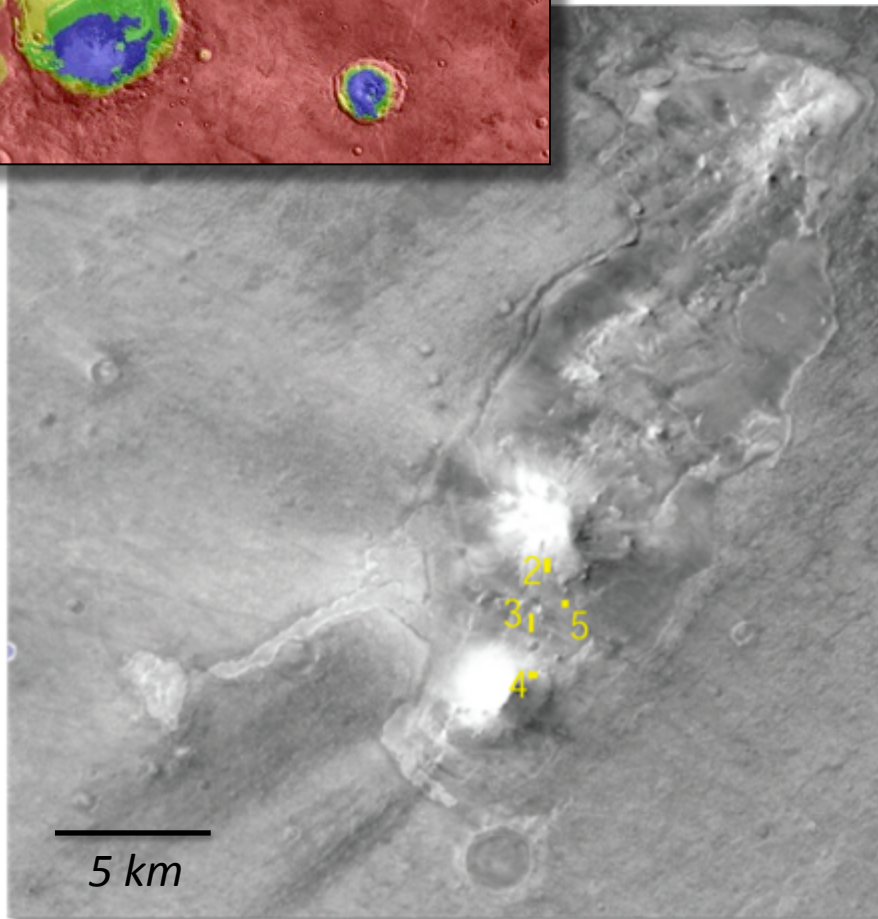
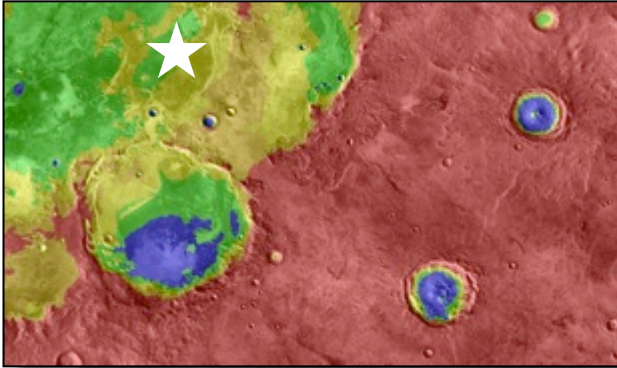
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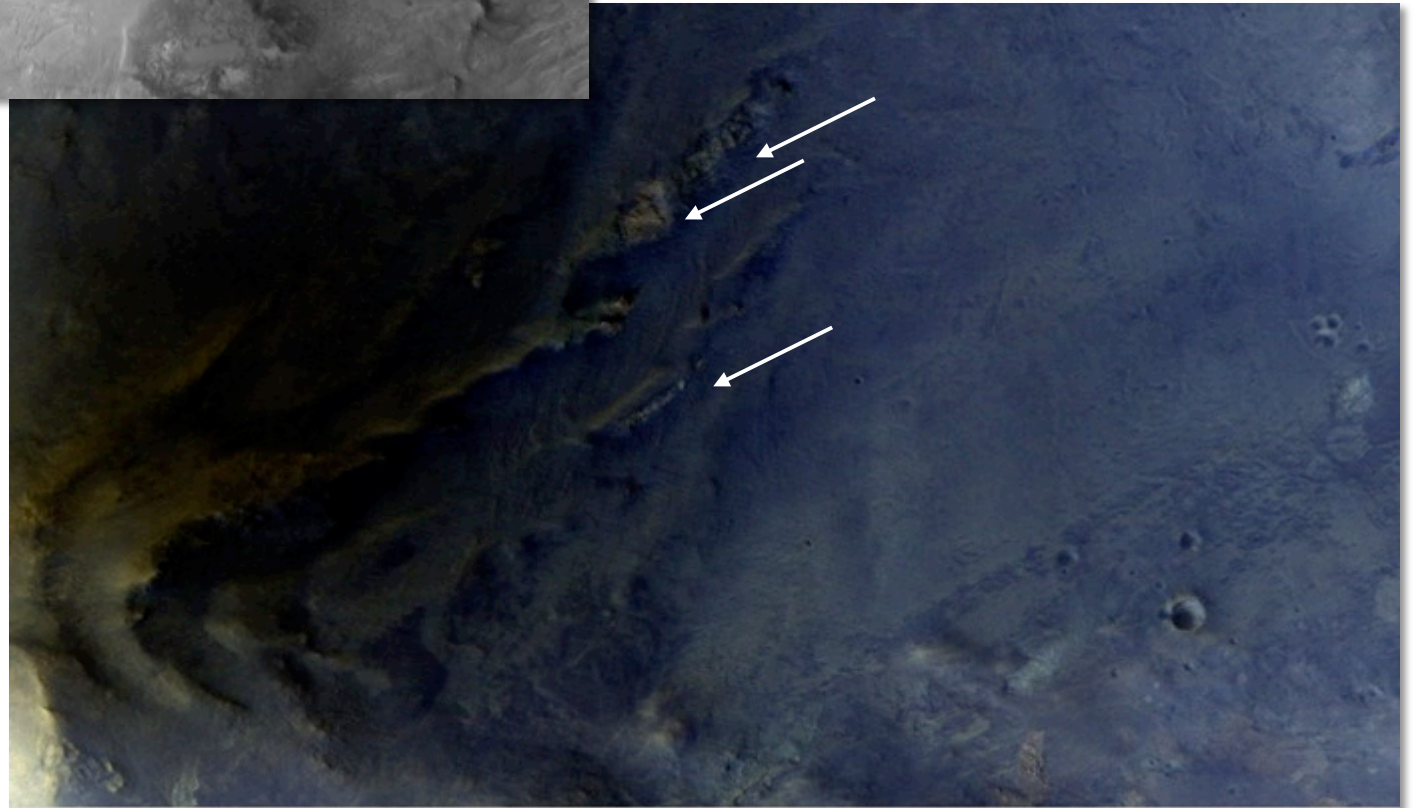
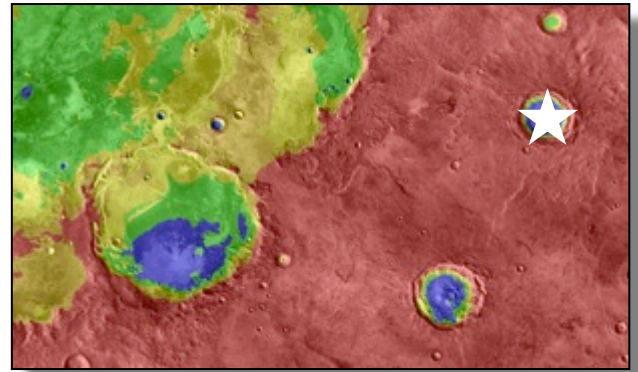


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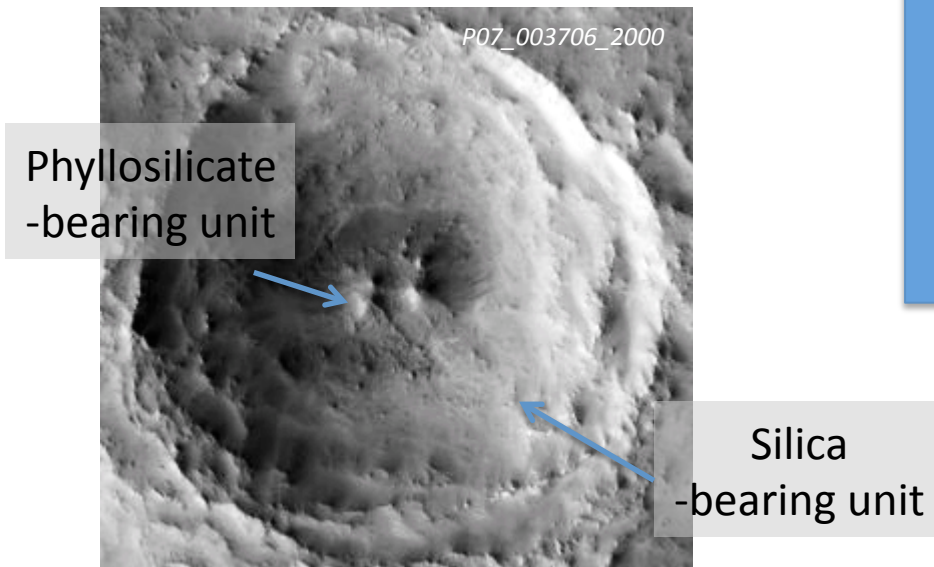
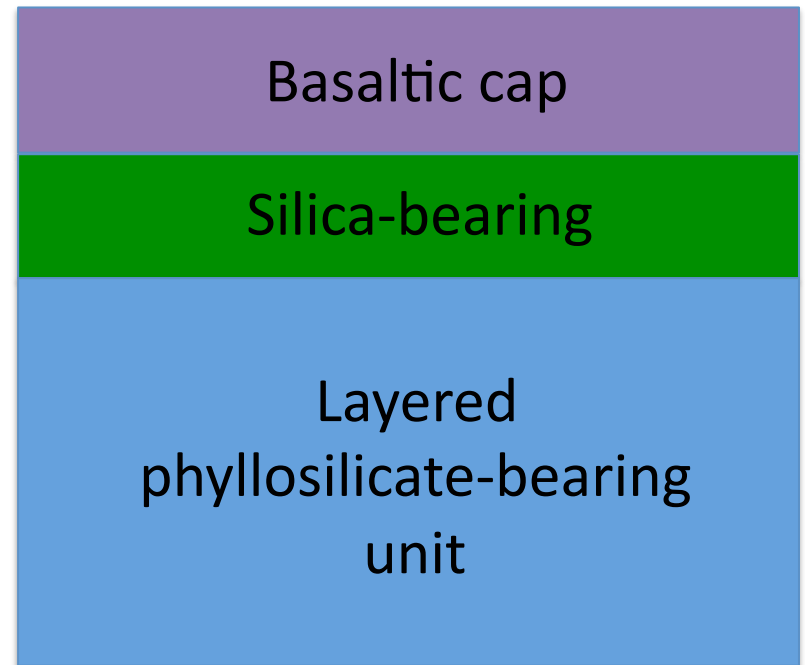
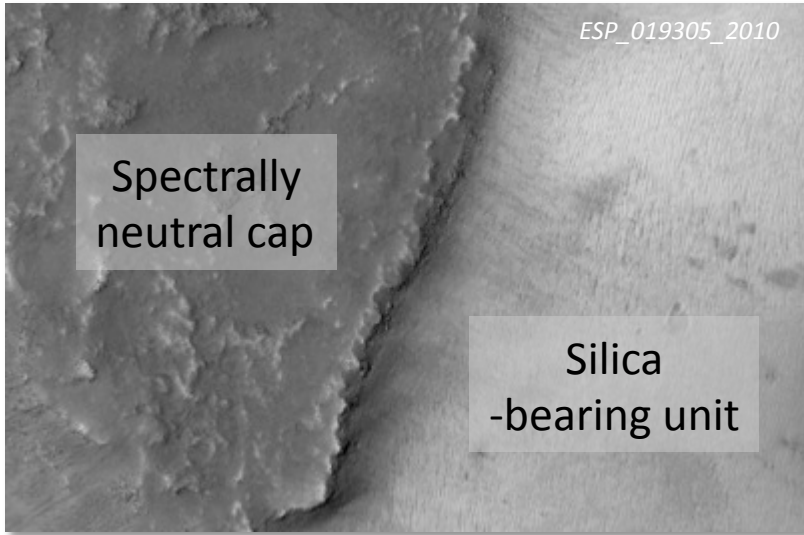
Layered knobs

PHYLLOSILICATES IN LAYERED KNOBS

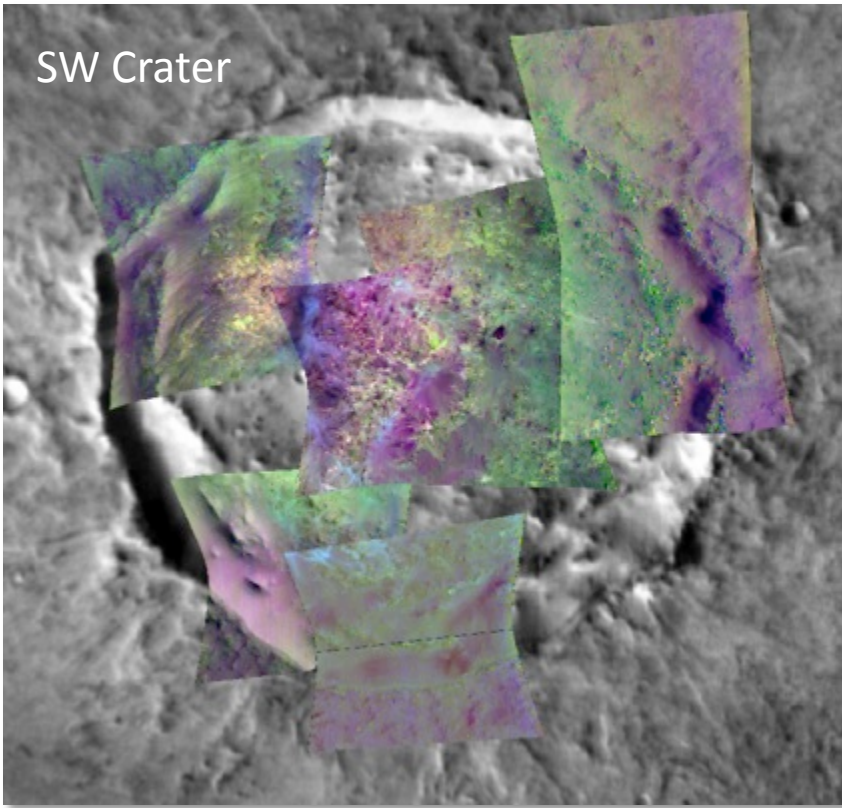




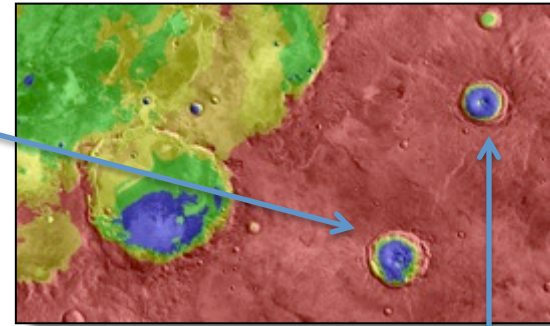
ANTONIADI CROSS-SECTION



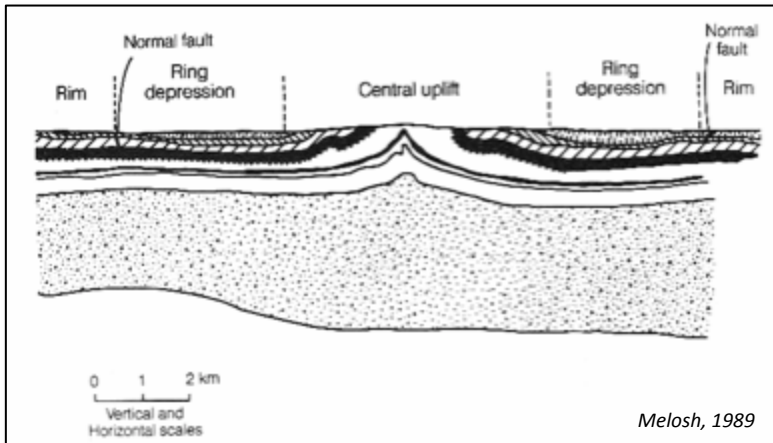
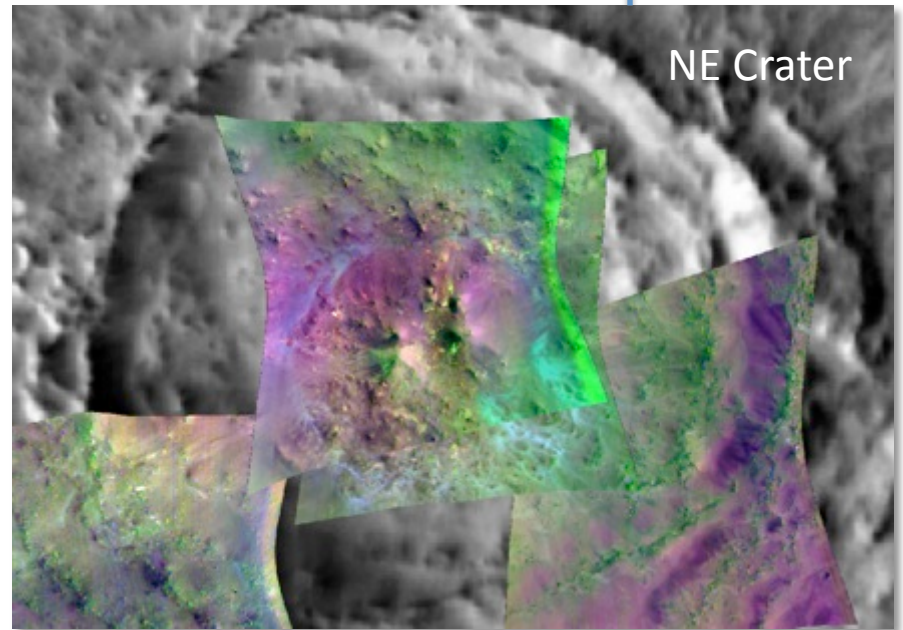
SW Crater



Blue = hydrated silica

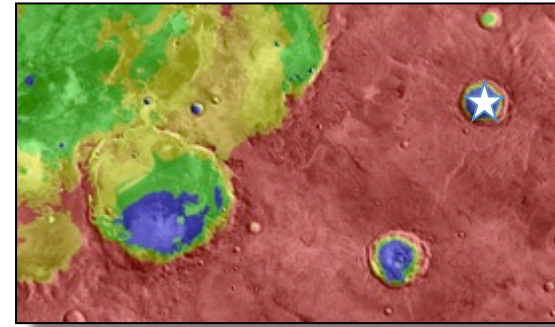
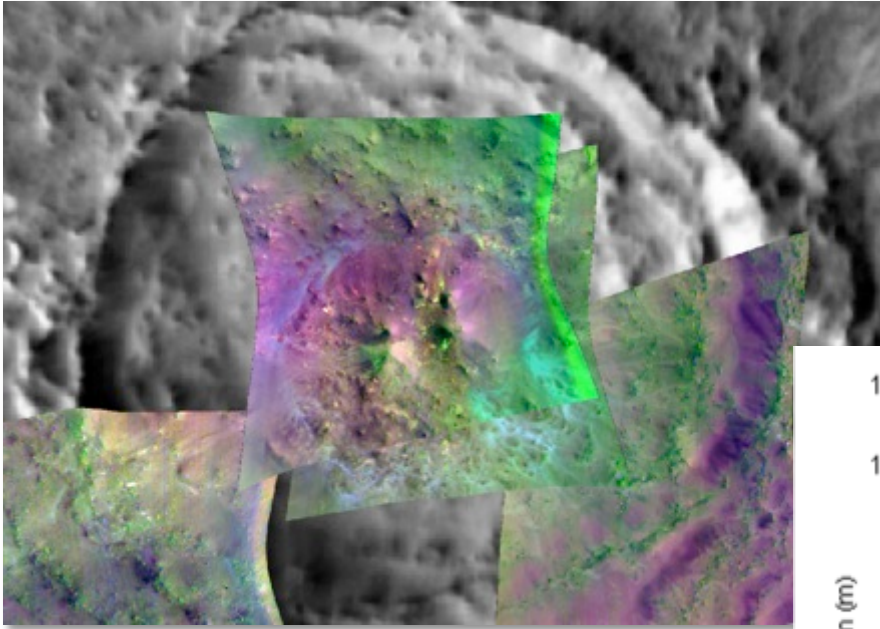


NE Crater



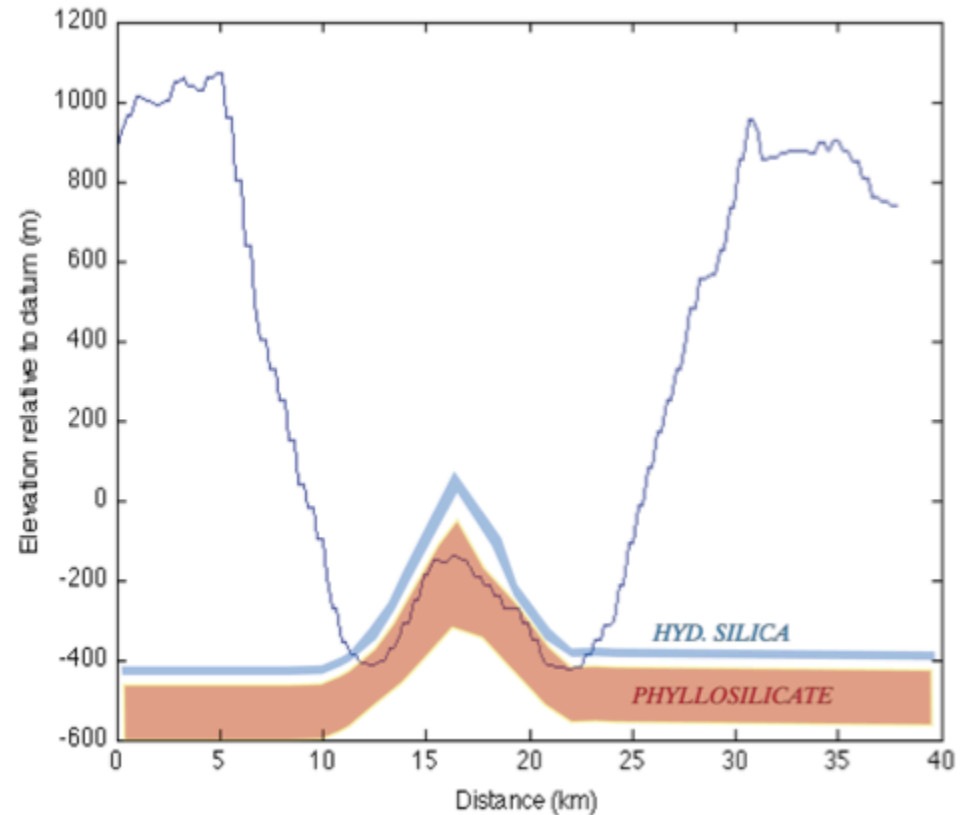
Centrally uplifted rocks are sourced from the deepest strata

NORTHEAST CRATER

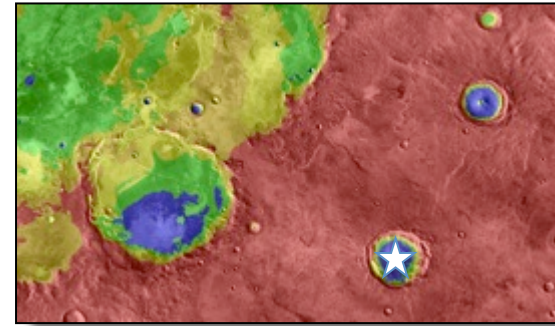
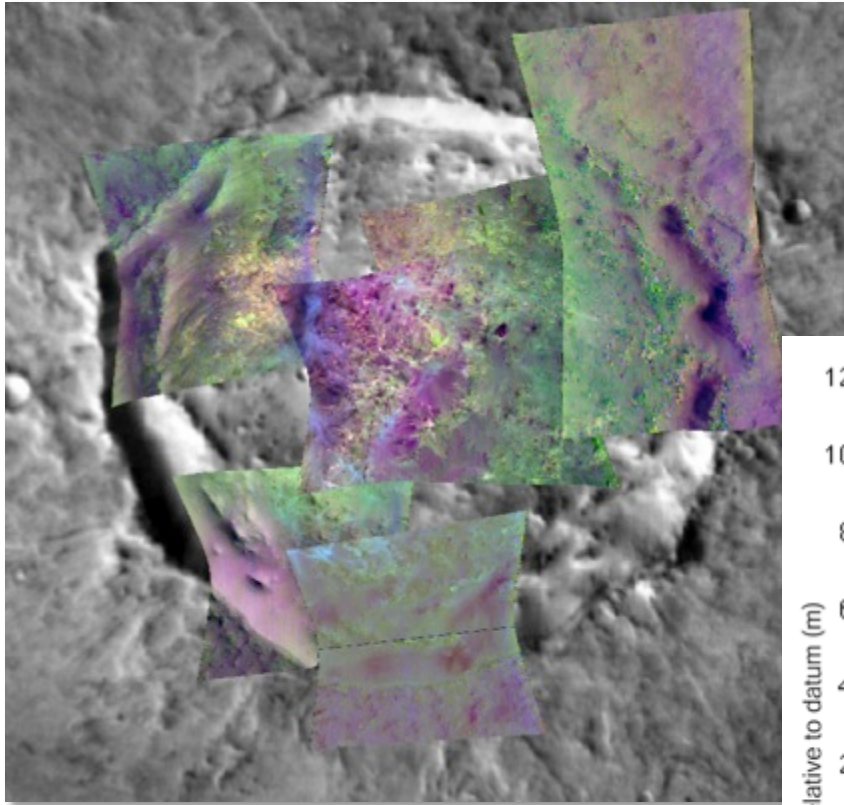


Hydrated silica is emerging from the walls, not from the central peak

Phyllosilicates exposed in the central peak are mostly detected in pre-existing layered terrain

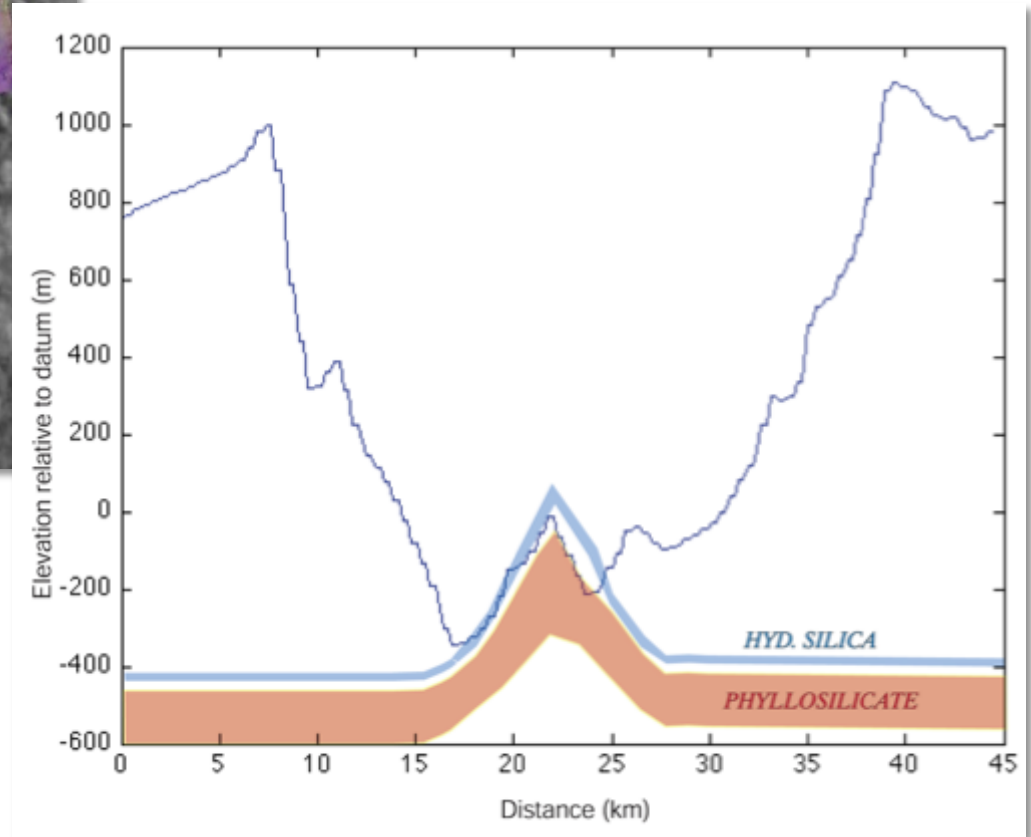


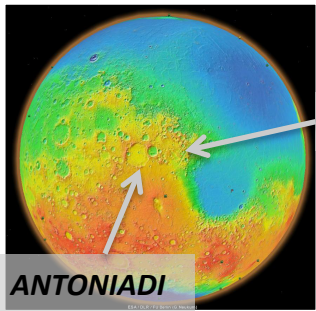
SOUTHWEST CRATER



Hydrated silica is eroding from the central peak, instead of the walls

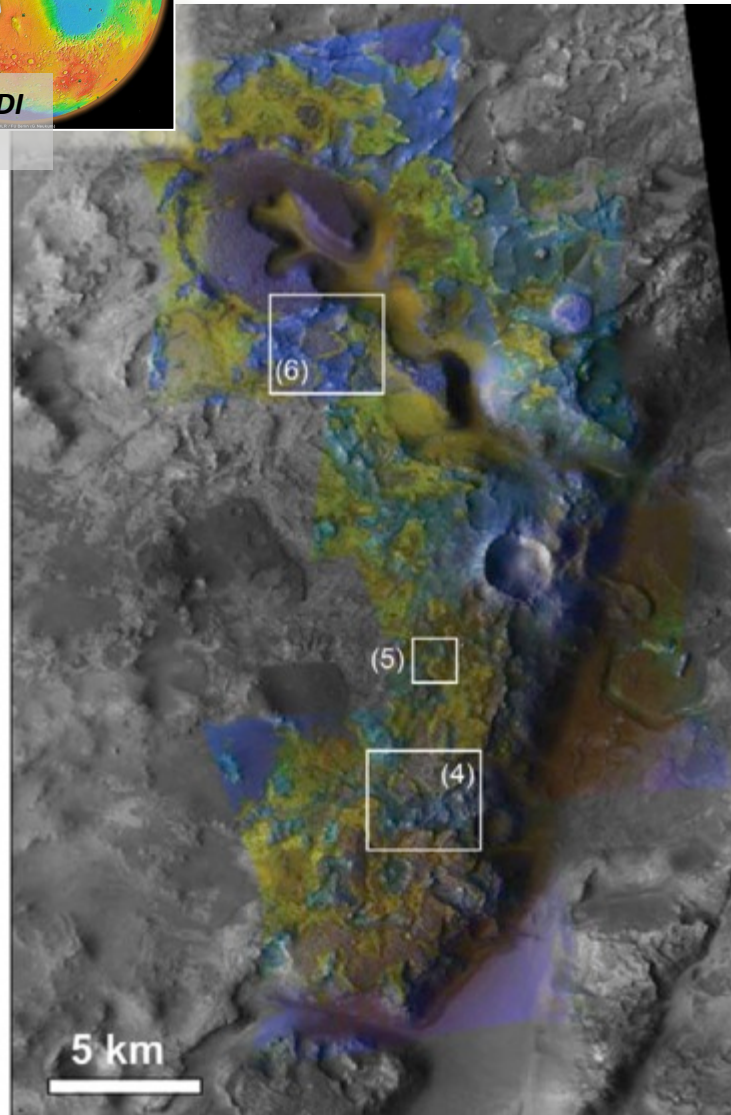
Phyllosilicates are found entirely in isolated blocks ringing the central peak, excavated during the impact process





NILI FOSSAE

ANTONIADI
CRATER



REGIONAL CONTEXT: *NILI FOSSAE*

Mafic cap

Olivine-rich basalt/carbonate unit

Kaolinite unit

Phyllosilicate-bearing
basement

*Olivine-bearing unit is inferred to be
impact melt from nearby Isidis impact*

Mustard et al., 2009

NILI FOSSAE AND ANTONIADI CRATER

- Pre-Isidis phyllosilicate-bearing basement is consistent between sites

However,

- *Isidis* contains an olivine/carbonate- and kaolinite-bearing layer, denoting a second and more extensive diagenetic event
- *Antoniadi* has a secondary-quartz-bearing unit, also suggesting extended periods of alteration

Differences in *protolith* composition can account for these differing mineralogical expressions of alteration

GLOBAL CONTEXT

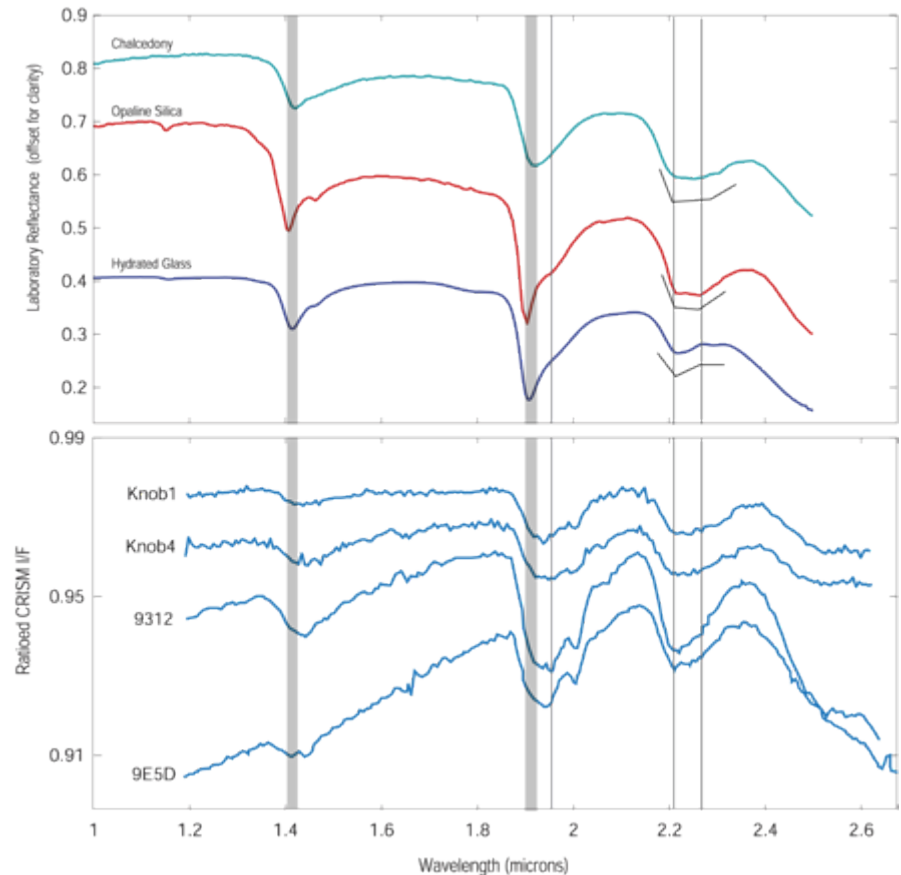
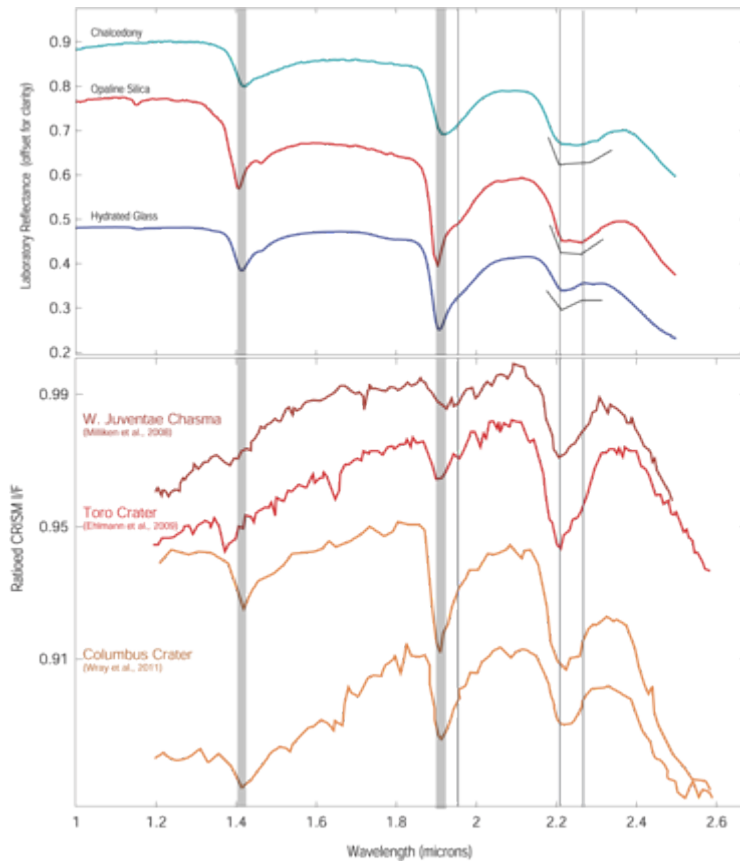
Unique detection of silica

- High-silica deposits form in two ways:
 - 1) Dissolution and reprecipitation of silica in an alkaline environment
 - 2) Dissolution of *everything else* in an acidic environment, leaving silica behind
- Hydrated silica is **not** associated with sulfates, unlike identifications at Valles Marineris (*Milliken et al., 2008*) and Gusev Crater (*Squyres et al., 2008*)
- Sulfate deposition indicates alteration in an acidic environment (*Formation #1*), clay deposition suggests alkaline (*Formation #2*)

GLOBAL CONTEXT

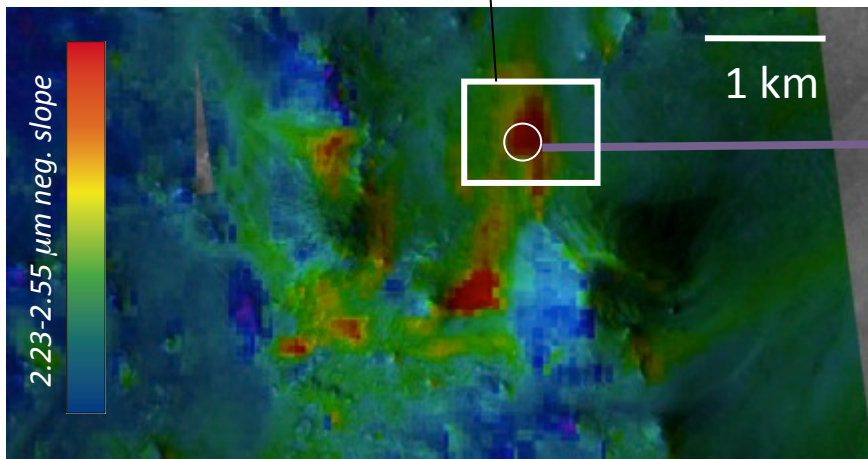
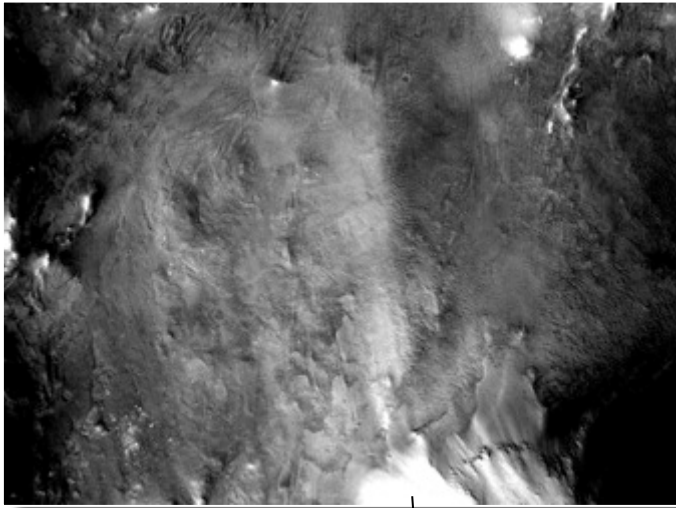
Unique detection of silica

- VNIR spectra show greater alteration than elsewhere on Mars

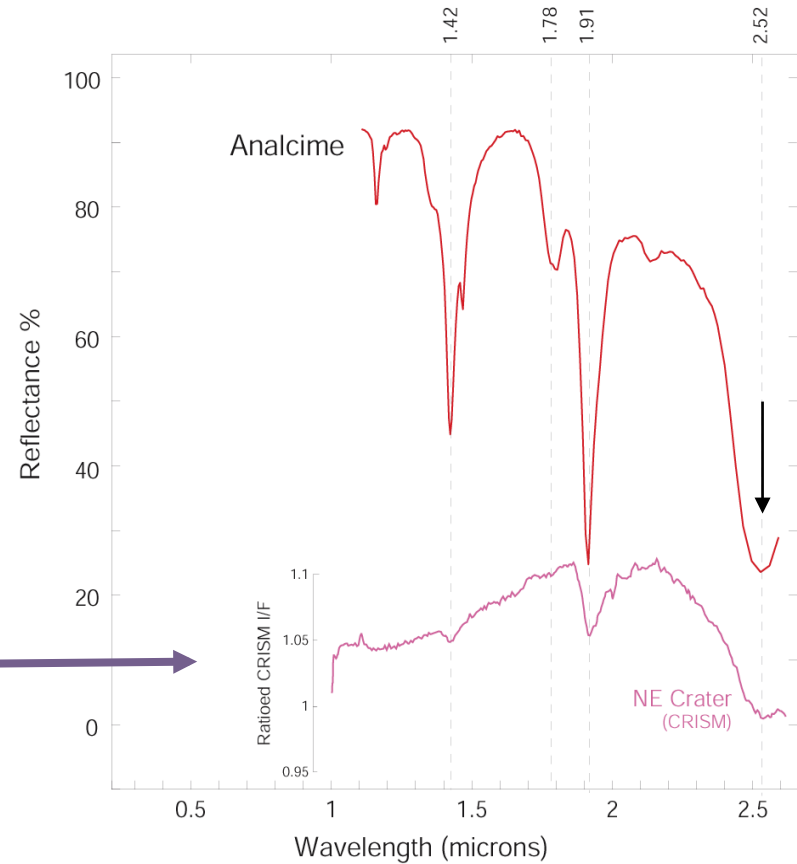


OUTSTANDING QUESTIONS

DO WE DETECT ZEOLITES?

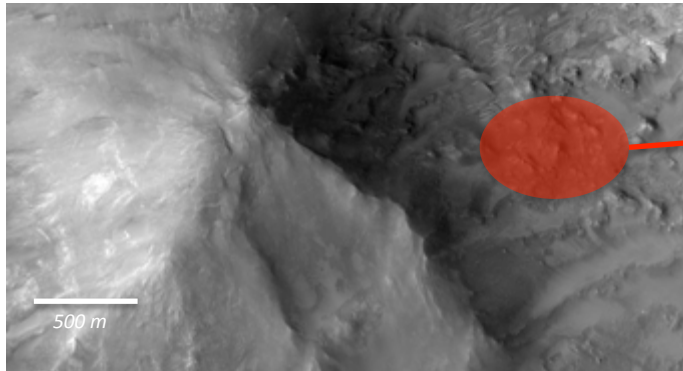
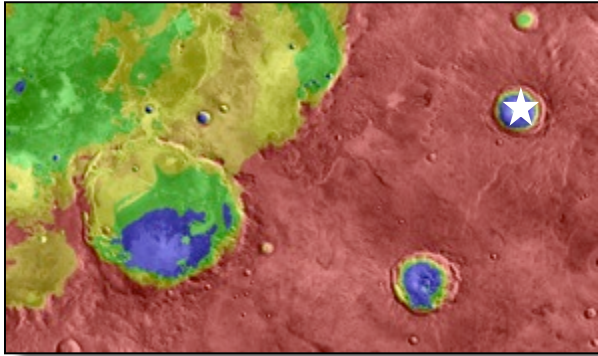


Zeolite detection

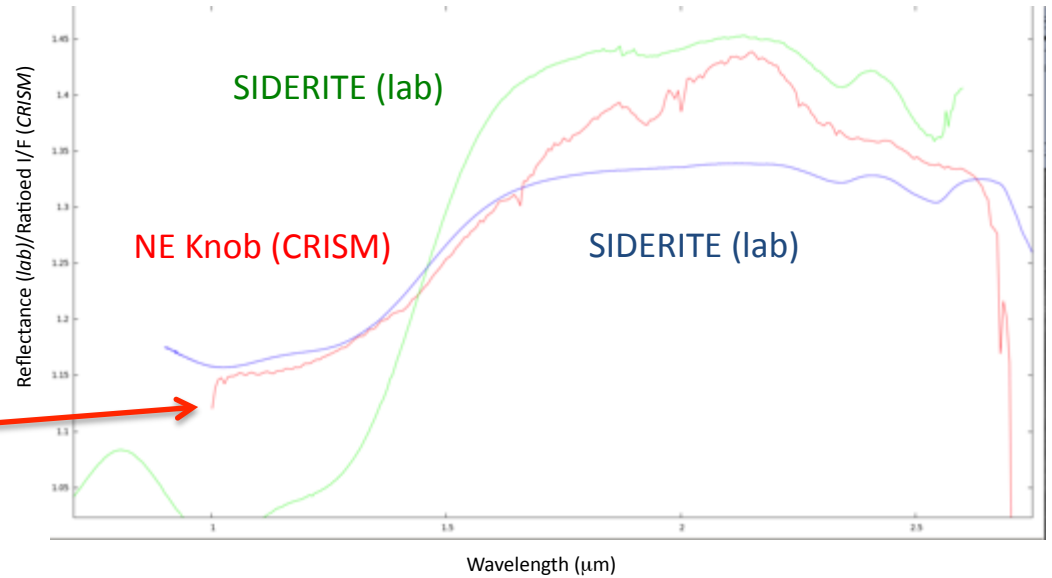


OUTSTANDING QUESTIONS

ARE WE ALSO DETECTING CARBONATES?



PSP_003706_2000



CONCLUSIONS

- Coincident with prior quartzofeldspathic detections, we detect chalcedony (in TES/CRISM) suggesting sustained local water-rock contact, and greater alteration than for other hydrated silica detections elsewhere on Mars
- We construct a regional stratigraphy with a Noachian phyllosilicate-bearing basement beneath a silica-bearing unit, suggesting multiple periods of wetting and alteration
- The stratigraphy is similar to nearby Nili Fossae, sharing both a phyllosilicate-bearing basement beneath a more altered layer suggesting multiple stages of wetting/alteration